

Comparison of Prototype and Laboratory Experiments on MOMA GCMS: Results from the AMASE11 Campaign

Sandra Siljeström,^{1,2} Caroline Freissinet,^{3,4} Fred Goesmann,⁵ Harald Steininger,⁵ Walter Goetz,⁵ Andrew Steele,² Hans Amundsen,⁶ and the AMASE11 Team

Abstract

The characterization of any organic molecules on Mars is a top-priority objective for the ExoMars European Space Agency–Russian Federal Space Agency joint mission. The main instrument for organic analysis on the ExoMars rover is the Mars Organic Molecule Analyzer (MOMA).

In preparation for the upcoming mission in 2018, different Mars analog samples are studied with MOMA and include samples collected during the Arctic Mars Analog Svalbard Expedition (AMASE) to Svalbard, Norway. In this paper, we present results obtained from two different Mars analog sites visited during AMASE11, Colletthøgda and Botniahelvøya. Measurements were performed on the samples during AMASE11 with a MOMA gas chromatograph (GC) prototype connected to a commercial mass spectrometer (MS) and later in home institutions with commercial pyrolysis-GCMS instruments. In addition, derivatization experiments were performed on the samples during AMASE11 and in the laboratory. Three different samples were studied from the Colletthøgda that included one evaporite and two carbonate-bearing samples. Only a single sample was studied from the Botniahelvøya site, a weathered basalt covered by a shiny surface consisting of manganese and iron oxides.

Organic molecules were detected in all four samples and included aromatics, long-chained hydrocarbons, amino acids, nucleobases, sugars, and carboxylic acids. Both pyrolysis and derivatization indicated the presence of extinct biota by the detection of carboxylic acids in the samples from Colletthøgda, while the presence of amino acids, nucleobases, carboxylic acids, and sugars indicated an active biota in the sample from Botniahelvøya.

The results obtained with the prototype flight model in the field coupled with repeat measurements with commercial instruments within the laboratory were reassuringly similar. This demonstrates the performance of the MOMA instrument and validates that the instrument will aid researchers in their efforts to answer fundamental questions regarding the speciation and possible source of organic content on Mars. **Key Words:** MOMA—ExoMars—Pyrolysis-GCMS—Derivatization—Svalbard—AMASE. *Astrobiology* 14, 780–797.

1. Introduction

THE SEARCH FOR SIGNS of past and present life on Mars is the main objective of the ExoMars 2018 mission, the European Space Agency (ESA) and Russian Federal Space Agency (Roscosmos) mission to Mars planned to launch in 2018 (Baglioni *et al.*, 2006). Though water has been detected on Mars, organic molecules, another important component of terrestrial life, still remain to be definitively identified *in situ*. The investigation of organics on Mars has so far been limited

to the biological and chemistry experiments of the Viking landers (Biemann *et al.*, 1977; Levin and Straat, 1977), the Sample Analysis at Mars (SAM) instrument on board the Curiosity rover (Mahaffy *et al.*, 2012), and investigation of martian meteorites.

The analysis of martian surface soils (upper 10 cm) by thermal volatilization–gas chromatography–mass spectrometry on the Viking landers did not detect any organic molecules other than two chlorinated hydrocarbons, which were believed at the time to be terrestrial contaminants (Biemann *et al.*, 1977;

¹Department of Chemistry, Materials and Surfaces, SP Technical Research Institute of Sweden, Borås, Sweden.

²Geophysical Laboratory, Carnegie Institution of Washington, Washington, DC, USA.

³NASA Goddard Space Flight Center, Greenbelt, Maryland, USA.

⁴NASA Postdoctoral Program (NPP), Oak Ridge Associated Universities, Oak Ridge, Tennessee, USA.

⁵Max-Planck-Institut für Sonnensystemforschung, Göttingen, Germany.

⁶Earth and Planetary Exploration Services, Oslo, Norway.

Biemann, 2007). This conclusion, however, has been challenged (Navarro-González *et al.*, 2006, 2010, 2011), defended (Biemann, 2007; Biemann and Bada, 2011), and debated (Navarro-González and McKay, 2011). The nondetection of organics by Viking was unexpected considering the model estimates of up to 500 ppm of organic molecules on the martian surface from meteorites, micrometeorites, and cometary infall (Flynn, 1996; Benner *et al.*, 2000). The lack of organics on the surface of Mars could be the result of their destruction by oxidants present in the soil and by the intense UV, solar, and cosmic radiation (Flynn, 1996; Benner *et al.*, 2000; Pavlov *et al.*, 2012). One of the biological experiments on the Viking landers, the “labeled release” experiment, had a positive response based on prelaunch criteria (Levin and Straat, 1977), which was interpreted by the designers of the experiment as evidence of life in martian soil (Levin, 1997). However, the nondetection of organics by the Viking gas chromatograph–mass spectrometer and the specifics of the release experiment led the Viking team to conclude that the labeled release was in response to an unknown inorganic oxidant in the sample (Biemann *et al.*, 1977). Later, the subsequent discovery of perchlorate in martian soil by the Phoenix mission (Hecht *et al.*, 2009) seemed to confirm this. Today, most agree that the Viking mission did not detect signs of life on Mars (Klein, 1999).

More recently, the Curiosity rover confirmed the presence of an oxychlorine phase in martian soil and rocks, most likely in the form of perchlorates (Glavin *et al.*, 2013; Leshin *et al.*, 2013; Ming *et al.*, 2014). The SAM instrument on board Curiosity has detected chlorinated organics (Glavin *et al.*, 2013; Leshin *et al.*, 2013; Ming *et al.*, 2014) that are thought to be derived from martian oxychlorine phases that interact with SAM baseline terrestrial carbon as well as a possible martian organic phase. A 100% terrestrial contribution to the formation of the detected chlorohydrocarbons has not been ruled out (Glavin *et al.*, 2013; Leshin *et al.*, 2013; Ming *et al.*, 2014), although with increased-fidelity measurements and quantitative background correction the possibility for a martian organic input appears to be more likely (Freissinet *et al.*, 2014).

An organic carbon component has been detected in martian meteorites (Wright *et al.*, 1989; McKay *et al.*, 1996) of which numerous sources have been proposed including contamination (Jull *et al.*, 1998; Stephan *et al.*, 2003), chondritic meteoritic input (Benner *et al.*, 2000), biology (McKay *et al.*, 1996), and various abiotic processes (Zolotov and Shock, 2000; Treiman, 2004; Steele *et al.*, 2007, 2012b). Recent *in situ* studies of the organic component in 11 martian meteorites concluded that a significant portion of this organic carbon is indigenous to Mars and that Mars has

its own internal source of organic carbon in the form of polycyclic aromatic hydrocarbons (Steele *et al.*, 2012a).

The detection of perchlorates at two very different locations on Mars (Hecht *et al.*, 2009; Glavin *et al.*, 2013) and in the martian meteorite EETA79001 (Kounaves *et al.*, 2014) suggests that distribution of oxychlorine phases is widespread at Mars’ surface and that pyrolysis analyses of martian surface samples will be challenging (Glavin *et al.*, 2013), as the chlorine present in these samples will interact with any organics during analyses (Steininger *et al.*, 2012). The ability of the ExoMars rover to retrieve samples 2 m into the regolith and sediments (Baglioni *et al.*, 2006) might mitigate some of these effects, as samples that contain more pristine organic molecules and less oxidants are more likely to be found below the surface, especially a few meters down, where they have been protected from solar and cosmic radiation (Dartnell *et al.*, 2007; Pavlov *et al.*, 2012).

The main instrument for organics analysis on ExoMars is the Mars Organic Molecule Analyzer (MOMA) (Goetz *et al.*, 2011a). MOMA has two fundamentally different operational modes: (a) pyrolysis–gas chromatography–mass spectrometry and (b) laser desorption ionization mass spectrometry. In short, MOMA consists of ovens where the sample will be deposited for single use, a gas chromatograph (GC) with four different column types, a pulsed UV laser, and a mass spectrometer (MS, designed as a 2-D ion trap).

The pyrolysis-GCMS experiment is capable of heating a sample to high temperature, $\sim 1000^{\circ}\text{C}$, and the gases evolved from the sample are then sent directly to the MS or to the GCMS unit for separation and identification of the components. To extend the range of detectable molecules to polar and nonvolatile compounds such as amino acids, carboxylic acids, and nucleic bases, MOMA has the capability of performing thermochemolysis and derivatization on the sample, as some of the MOMA ovens will contain capsules filled with tetramethylammonium hydroxide (TMAH) for thermochemolysis and *N,N*-methyl-*tert*-butyl-dimethylsilyltrifluoroacetamide (MTBSTFA) or dimethylformamide dimethylacetal (DMF-DMA) for derivatization (Rosenbauer *et al.*, 1999). In addition, DMF-DMA allows derivatization with conservation of enantiomeric centers of the molecules, enabling the chiral analysis of a wide range of exobiology-relevant refractory organic molecules. Derivatization, which is complementary to pyrolysis, allows analysis of more complex molecules due to two intrinsic principles. First, derivatization makes the molecules more volatile by capping any reactive groups and thus prevents them from interacting with tubes and the chromatographic column (Fig. 1). Polar organic molecules such as amino acids and carboxylic acids are too refractory for



FIG. 1. The derivatization reaction with MTBSTFA on an amino acid. The reactive groups (labile H) are silylated by the derivatization reagent, protecting the molecule and making it more volatile and thus amenable to gas chromatography–mass spectrometry. The product is a tBDMs ester of the initial molecule. Color images available online at www.liebertonline.com/ast

normal pyrolysis–gas chromatography–mass spectrometry. Second, contrary to pyrolysis, which breaks down the heavy components, derivatization protects the high-molecular-weight organic molecules. In fact, it has been shown (Glavin and Bada, 1998) that amino acids require chemical extraction from the solid mineral matrix prior to pyrolysis heating to avoid such decomposition. Moreover, the reaction occurs at mild temperatures, which helps to preserve the integrity of the molecules of interest. However, this chemical transformation also adds complexity to the analysis of derivatized compounds.

In preparation for missions to Mars, martian analogues on Earth are investigated with flightlike protocols and instrumentation (Steele *et al.*, 2006; Bishop *et al.*, 2013). Samples from these types of environments typically contain low levels of organic and biological material in mineralogical assemblages relevant to Mars (Navarro-González *et al.*, 2003, 2006). The Svalbard archipelago, situated in the Arctic Ocean north of mainland Europe (74° to 81°N, 10° to 35°E), has a cold (average temperature -6°C) and dry (210 mm of annual rainfall in Longyearbyen) climate and harbors numerous sites analogous to Mars (Steele *et al.*, 2006). In addition, minimal vegetation and the presence of glacier valleys that cut through the sedimentary layers has resulted in excellent rock exposure and easy access to numerous sites where different aspects of martian habitability might be explored (Steele *et al.*, 2006; Blake, 2010). Svalbard has been visited since 2003 by the Arctic Mars Analog Svalbard Expedition (AMASE), where a team of scientists and engineers who study the Mars analog environment use instruments and techniques that may be used for current and future planetary missions, including the Mars Science Laboratory, ExoMars, and Mars Sample Return (Steele *et al.*, 2006). During AMASE 2010 and AMASE 2011, the MOMA team employed a GC breadboard to test the MOMA prototype instrument on samples in a Mars analog environment. These samples were then retrieved from the field and investigated in home institutions with commercial instrumentation under more optimal conditions. This paper describes analysis results of samples from two different field sites, Colletthøgda and Botniahalvøya, which were visited by way of the research vessel (R/V) Lance during AMASE11. This was the first successful coupling of a MOMA-type GC system with a MS for field analysis of samples. Because the MOMA MS was under development and not ready for field testing when these experiments were performed, a commercial MS had to be used. All the reagents used and experiments performed in this study met the technological and operational constraints imposed by spaceflight experiments; thus these analyses demonstrated the capabilities of near-term landed missions.

2. Field Sites on Svalbard

During the AMASE11 campaign, two different sites on Svalbard, Colletthøgda and Botniahalvøya, were visited and found to be good Mars analogues with a range of relevant mineralogies that include evaporites, carbonates, and weathered basalts (Fig. 2).

2.1. Colletthøgda

Colletthøgda is situated on the west coast of Svalbard just north of Ny Ålesund (Fig. 2A, 78.89°N, 12.55°E) and is a

marine succession that is part of the Carboniferous/Permian Dickson land subgroup and the superior Gipsdalen group, which cover parts of central Svalbard (Dallmann, 1999). The Gipsdalen group represents the development from a clastic continental graben to an open-shallow marine shelf with the Dickson land subgroup representing the later part of this development (Dallmann, 1999, 2002). The Dickson land subgroup consists of mainly carbonate rocks, represented by the Wordiekammen Formation, with subordinate evaporites, the Gipshuken Formation (Dallmann, 1999). At the field site at Colletthøgda (Fig. 2A), the Wordiekammen Formation is exposed at the bottom as layered strata, which the Gipshuken Formation overlies as a homogeneous structure. In addition, at Colletthøgda another homogeneous structure can be seen intersecting the Gipshuken and Wordiekammen Formations (Fig. 2A). It has been proposed that this structure is a bioherm, that is, a reeflike structure.

Features that contain similar mineralogy to that found at Colletthøgda are found on Mars. Evaporites have been identified in martian meteorites such as Nakhla and EETA 79001 (Gooding, 1992; Sawyer *et al.*, 2000) and in various places on Mars (Squyres *et al.*, 2004; Gendrin *et al.*, 2005; Langevin *et al.*, 2005; Bibring *et al.*, 2006) including Gale Crater (Milliken *et al.*, 2010). Evaporites have been suggested to be a good habitat for microbes in an extreme environment such as the surface of Mars (Rothschild, 1990) and have shown to have good preservation potential for organic matter (Aubrey *et al.*, 2006). Carbonates have been found on Mars (Ehlmann *et al.*, 2008; Morris *et al.*, 2010) and in martian meteorites (McKay *et al.*, 1996). Although a bioherm is not expected to be found on Mars, it is a suitable analogue for biologically produced carbonates.

2.2. Botniahalvøya

Botniahalvøya is situated on the Nordaustlandet island in the northeast of the Svalbard archipelago (Fig. 2B, 80.25°N, 19.25°E). At the field site, Precambrian lava (the Kaap Hansteen suite) is cut by dolerite sills that are part of the Diabasodden suite (Fig. 2B–2C) (Dallmann, 2002). This suite is found mainly in parts of eastern Svalbard and is believed to have formed during the early Cretaceous (Dallmann, 1999, 2002). At the field site on Botniahalvøya, the sill is exposed and has been weathered by water seeping through it, which has created a horizon of differentially weathered basalts. Some parts of the weathered surface have a shiny appearance much like desert varnish. Desert varnish is a type of coating that consists of clay minerals cemented to rock by oxyhydroxides of manganese and iron, which are found on many rocks in arid environments (Dorn, 1998). It has been suggested that biology is involved in the formation of desert varnish (Dorn and Oberlander, 1981; Taylor-George *et al.*, 1983).

The rocks in the field site at Botniahalvøya are analogues to weathered basalts found on Mars. During the Viking mission, rocks with shiny surfaces were reported that could possibly be desert varnish (Moore *et al.*, 1987; Arvidson *et al.*, 1989; Guinness *et al.*, 1997). It has been suggested that desert varnish could be a potential habitat for martian microorganisms (DiGregorio, 2002). Although recent findings of desert varnish at Gale Crater are not indicators of biology *per se*, this exciting find suggests that further investigation of

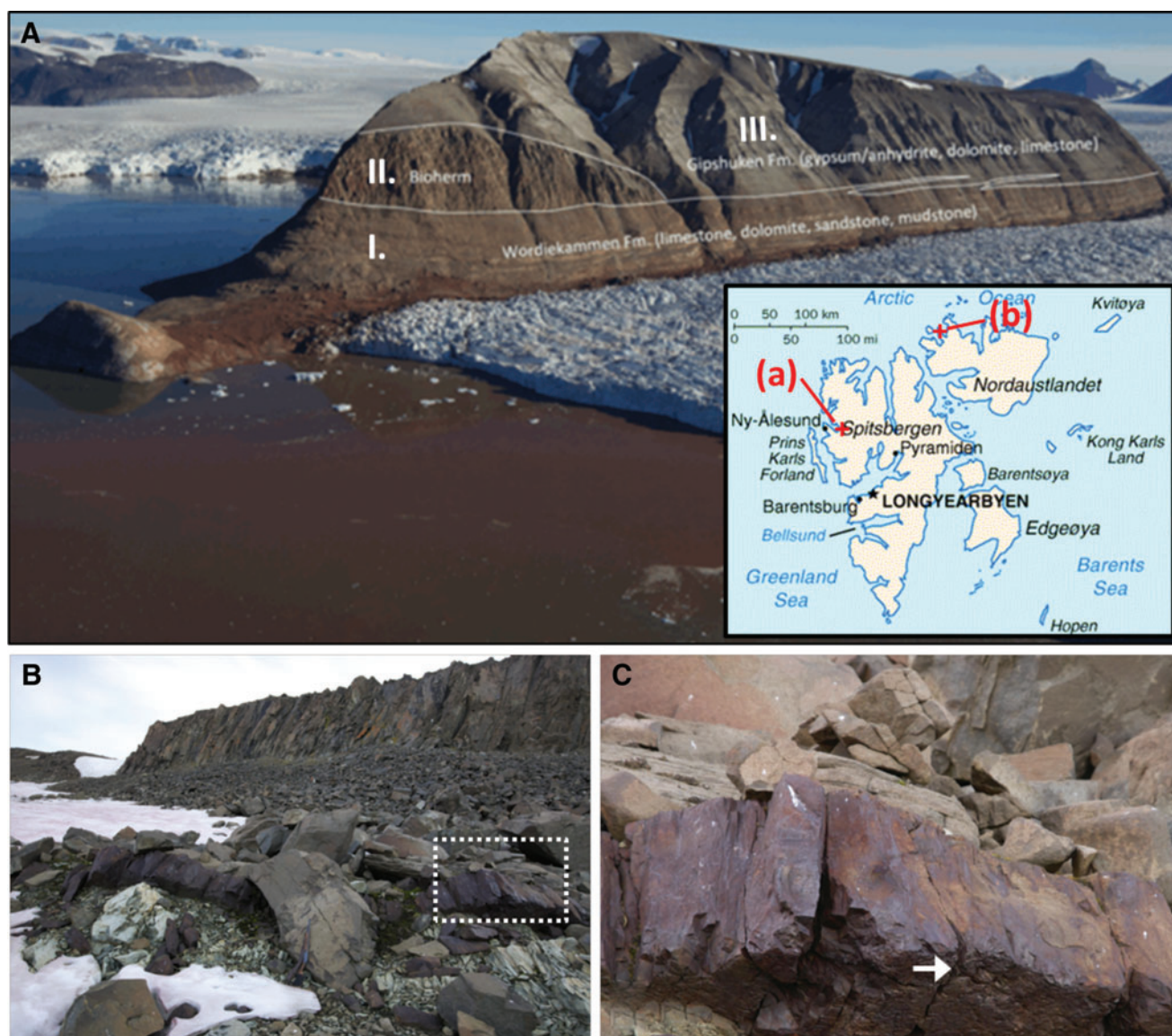


FIG. 2. Field sites on Svalbard. (A) Colletthøgda and (B) Botniahalvøya. Arrows in the map (inset) indicate the locations of the two field sites. Formations I, II, and III in (A) are, respectively, the Wordiekammen Formation, a bioherm, and the Gipshuken Formation. White rectangle in (B) indicates sampled sill in (C). White arrow in (C) indicates site of sample AM11-162. Photo credit: Hans Amundsen, Kjell Ove Storvik, and the CLUPI team. Color images available online at www.liebertonline.com/ast

terrestrial analogues would be advantageous (Lanza *et al.*, 2014). In addition, various clay assemblages have been found on Mars (Bibring *et al.*, 2006; Vaniman *et al.*, 2014), and hematite is known to be present in the martian regolith (Ming *et al.*, 2006; Morris *et al.*, 2006; Vaniman *et al.*, 2014).

3. Materials and Methods

3.1. Samples

Three different samples (Table 1) were studied from the Colletthøgda site: an evaporite (Snickers) and two carbonates (KitKat and Butterfinger). One sample was studied from the Botniahalvøya site, AM11-162 (Table 1).

The samples from Colletthøgda (KitKat, Butterfinger, and Snickers) were collected as float at the base of the outcrop at Colletthøgda, while the Botniahalvøya sample was collected

from the exposed outcrop. The collected samples were immediately wrapped in ashed aluminum foil before being placed in a labeled plastic bag. The rocks were analyzed in laboratories on board the R/V Lance and were initially cracked open with a rock hammer. Small pieces of the freshly fractured rock were collected from the inner part (Colletthøgda) or from the outer part (Botniahalvøya) of the samples. The collected pieces of rocks were ground to powder with an agate mortar and pestle previously cleaned with distilled water, isopropanol, and ethanol (Eigenbrode *et al.*, 2009).

3.2. X-ray diffraction/X-ray fluorescence

X-ray diffraction (XRD)/X-ray fluorescence (XRF) measurements of the bulk mineralogy of the samples were performed with the Terra XRD/XRF (InXitu Inc., Mountain

TABLE 1. SAMPLES ANALYZED AND THEIR MINERALOGICAL COMPOSITION

Samples	Sample description	Terra XRD
Colletthøgda (Snickers)	White rock, well rounded with anastomosing (gray and black veins).	Mainly anhydrite + 9% gypsum
Colletthøgda (KitKat)	Buff-colored carbonate mudstone.	Mainly dolomite + some quartz
Colletthøgda (Butterfinger)	Network of subparallel veins with a diameter of 0.25 cm homogeneously distributed. Possible carbonate mud between veins.	Largely calcite, some dolomite and quartz
Botniahelvøya (AM11-162)	Underside of the sill. Dark and greasy.	Least-altered part of sill: albite 37.5%, diopside 23.2%, quartz 6.4%, illite 23.2%, chrysotile 5.3%, ilmenite 4.4% Dark greasy sill coating: albite 13.1%, diopside 7.9%, quartz 10.6%, ramsdellite 4.8% (MnO), smectite (RIR = 3) 54.0%, hematite 9.6%

View, USA). The Terra instrument is similar to the Chemistry and Mineralogy (CheMin) instrument on board the NASA Curiosity rover (Blake, 2010).

3.3. Pyrolysis–gas chromatography–mass spectrometry

3.3.1. Analyses during AMASE11

Pyrolysis. The hardware used during AMASE11 consisted of a simplified MOMA GC from the Max Planck Institute for Solar System Research (MPS) and a commercial GCMS (Griffin Minotaur 400 field GCMS) from NASA Goddard Space Flight Center (GSFC).

The MOMA GC hardware was based on the current MOMA flight hardware developments at MPS. It consists of MOMA ovens (stepped or ramped heating up to $\sim 1000^{\circ}\text{C}$), a MOMA tapping station interfaced to the ovens, a thermal injection trap (Tenax TA) where evolved volatiles are trapped at temperatures below 0°C , an analytical column module with a MXT20 capillary column ($20\text{ m} \times 0.25\text{ mm} \times 0.25\text{ }\mu\text{m}$; Restek Cooperation), a thermal conductivity detector (Varian Inc.), as well as prototype electronics controlled by in-house-developed LabView software.

A commercial Griffin Minotaur 400 field GCMS was used both as stand-alone for the derivatization experiments and as

the ion trap MS of the MOMA GC to perform MOMA-like pyrolysis–gas chromatography–mass spectrometry. The Griffin Minotaur 400 is designed for operation outside a typical laboratory environment, with larger ambient temperature range and relative humidity constraints. It contains an electron impact ionization source and a cylindrical ion trap mass analyzer. The MS assembly is operated at a pressure of 10^{-3} to 10^{-4} mbar, which is significantly higher than the pressure used for many other types of mass analyzers. The mass range of the Minotaur 400 is from 10 to 425 u.

During the AMASE11 field trip, the team managed to operate the two instruments (MOMA GC and the Griffin Minotaur GCMS) separately and successfully when interfacing the two together (Fig. 3). To interface the MOMA GC with the Griffin Minotaur MS, we by-passed the GC part of the Minotaur and connected a helium gas tank directly to the MOMA GC setup, which was further connected downstream to the Griffin MS (Fig. 3). This scheme allowed the MS and the GC to operate together in GCMS mode for separation (MOMA GC) and definitive identification (Griffin MS) of organic compounds, thus making it possible to obtain pyrolysis-GCMS data from the crushed rock samples during AMASE11.

The operating condition of the interfaced MOMA GC and Griffin Minotaur MS were as follows: the GC tubing and the

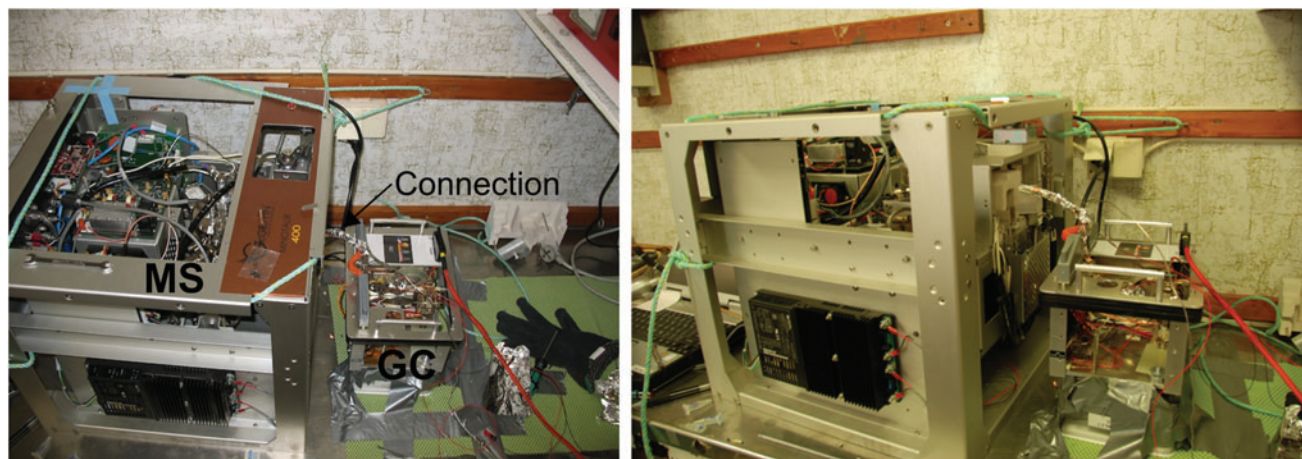


FIG. 3. Instrumental setup of MOMA GC to Griffin MS on board the R/V Lance during AMASE11. Color images available online at www.liebertonline.com/ast

GC column were preheated to 60–70°C and 30°C, respectively, and a constant gas pressure corresponding to ~2 mL/min He flow through the GC column was established. Around 200 µg of powdered rock sample (Butterfinger, KitKat, Snickers, or Botniahavøya) was placed in the oven. The oven was then heated to 600°C under He flow to pyrolyze and release any organics from the sample. The evolved gases were flushed and collected into the trap that was cooled down to <5°C. The trap was then heated to 250°C for around 10 s to release the volatiles and send them to the GC column. After injection, the column temperature was ramped at 5°C/min to 150°C. After elution from the GC column, the discrete molecules entered the thermal conductivity detector, where they were recorded. These molecules were then routed via the exhaust of the MOMA GC and a heated (60–70°C) tube into the ion trap MS of the Griffin Minotaur, which was operated independently. The MS was operated between 40 and 425 u to exclude signals from the air such as nitrogen and oxygen and thus improve the signal-to-noise ratio. Between each sample measurement, a complete measurement cycle as described above was carried out by using an empty oven to retrieve the baseline and keep track of contamination from the surrounding environment.

Derivatization. In this study, MTBSTFA was used as a derivatization reagent and dimethylformamide (DMF) as its solvent, which allowed an efficient and nonselective derivatization on a wide range of organic molecules. MTBSTFA performs silylation of any labile hydrogen and thus transforms polar organic compounds of low volatility, such as amino acids and carboxylic acids, into volatile *tert*-butyldimethylsilyl (tBDMS) derivatives that can be separated and identified by gas chromatography–mass spectrometry (Knapp, 1979).

To perform derivatization, the Griffin GCMS was used on a stand-alone base. The preparation of the sample included a deposit of a few milligrams (~10–20 mg) of the sample into a glass vial, the addition of 15 µL of MTBSTFA/DMF (4:1), the heating of this mixture during 10 min at 175°C in a heater block to allow both extraction of the organic content by thermodesorption and by the DMF solvent, and subsequent derivatization of the extracted organics (one-pot one-step procedure), followed by the direct injection of 2 µL of this sample into the GCMS. The sample was introduced into the GC inlet by direct injection of a liquid phase with a microsyringe, where it was rapidly heated and swept onto the analytical column by the He carrier. Only the samples from Colletthøgda were analyzed on board the R/V Lance.

In addition, because MTBSTFA leads to a lot of decomposition products that can interfere with the peaks of interest, and to ensure that the reaction is working, a blank was prepared by heating a solution of 10 µL of MTBSTFA/DMF (4:1) to 75°C during 10 min. Because of the constraints of an on-site experiment, no other blanks were run between the samples. Both blank and samples were run with a GC program from 70°C held 5 min to 300°C held 5 min, with a 10°C/min ramp, in splitless mode and with a constant He inlet pressure of 690 mbar. The operational mass range of the MS was between 40 and 425 u.

The Griffin MS has an unresolved offset of mass calibration of 1–2 u, which means that the NIST mass spectral

library cannot directly be used to identify the detected compounds. Therefore, commercial laboratory runs were performed to confirm whether a specific compound was present in the AMASE11 runs.

3.3.2. Laboratory analyses at home institutions

Pyrolysis. Pyrolysis was performed at MPS. A commercial Varian GCMS (GCMS 4000) coupled to a commercial pyrolysis unit (Pyrola) was used. A chromatographic column with the identical retention phase (MXT20: 20 m × 0.25 mm × 0.25 µm) as the MOMA GC was used. The temperature profiles of the oven and column reproduced those used during AMASE11, except for the Botniahavøya sample (AM11-162), which was pyrolyzed at 900°C instead of 600°C.

Derivatization. Derivatization experiments were performed at GSFC with a commercial Thermo Finnigan Trace GC Ultra DSQII quadrupole MS. It consisted of a solvent extraction of the organic molecules from the sample followed by a derivatization of the extract. The four samples were weighed in cleaned glass ampules: 187.2 mg of KitKat, 163.0 mg of Snickers, 81.5 mg of Butterfinger, and 149.2 mg of Botniahavøya. One milliliter of a mixture of H₂O/isopropanol (1:1) was then added to each sample. Blanks consisted of an ampule only filled with the H₂O/isopropanol (1:1) mixture. All the ampules were then sealed and placed in an oven at 100°C. After 24 h, the supernatant was transferred out to a clean vial where it was evaporated under vacuum. Five microliters of a mixture of MTBSTFA/DMF (4:1) was added, and the vials were then heated to 75°C for 15 min to allow the derivatization reaction to proceed, after which 0.1 µL of this solution was injected in the GCMS. The GCMS was set in splitless mode with a temperature program from 50°C to 300°C at 10°C/min and a 3 min hold at 300°C, under a constant pressure of 903 mbar of He. The MS scans from 3 to 535 u from 6 min of run to the end (solvent delay).

3.4. Scanning electron microscopy/energy dispersive X-ray analysis

The sample from Botniahavøya (AM11-162) was also investigated in a scanning electron microscope (SEM; Leitz Amray 1600T) with energy-dispersive X-ray (EDX; Tracor Northern) analysis to study the difference in morphology and chemistry between the outer (most altered) and inner (least altered) part of the sample. The samples were sputter-coated with 10 nm gold before analysis. SEM imaging was conducted at 30 kV with subsequent EDX spectra acquired with a similar acceleration voltage.

4. Results

4.1. Field XRD/XRF data

The data collected by Terra XRD/XRF during AMASE11 is summarized in Table 1. The mineralogy of the Butterfinger, KitKat, and Snickers samples is consistent with their respective assignment to the formations of Wordiekammen and Gipshuken as described in the literature (Dallmann, 1999, 2002). The mineralogy of Botniahavøya samples is that of a weathered basalt (Nejbert *et al.*, 2011) containing clay as well as iron and manganese oxides.

4.2. Pyrolysis and derivatization

4.2.1. Comparison between the AMASE11 prototype and commercial laboratory measurements

Pyrolysis. For direct comparison of the results from the MOMA prototype and commercial laboratory systems, total ion chromatograms (TICs) were obtained for the Butterfinger and Botniahalvøya (AM11-162) samples. Two sets of pyrolysis data were acquired for each set of samples: one set on the MOMA prototype during AMASE11 and the other on the commercial systems at MPS (Fig. 4, Tables 2 and 3). The chromatograms of Butterfinger from both systems (Fig. 4A) show the presence of numerous organic molecules including long-chained hydrocarbons and aromatic molecules such as benzene, naphthalene, and a range of alkylated benzenes, while the chromatograms of Botniahalvøya (AM11-162; Fig. 4B) show long-chained hydrocarbons, nitrogen-, oxygen-, and sulfur-containing aromatics, and

various kinds of heterocyclic compounds. In Fig. 4C, the mass spectra of the trimethylbenzene (peak 8 in Fig. 4A) found in the Butterfinger sample are compared. The mass spectra compared were retrieved from the Griffin MS, the commercial Varian system, and the NIST library. The data quality acquired on the MOMA prototype system was surprisingly good considering that this was the first joint operation of the MOMA GC and a commercial MS and the analyses took place in sub-optimal conditions on board the R/V Lance. The produced data were reproducible across several analyses, with different samples producing unique signatures that were different from the blanks analyzed during AMASE11. Any differences in results between the commercial and the prototype systems are explained by the fact that the operational parameters were optimized for the Svalbard setup, not for the commercial equipment. At low retention times (<5 min), the commercial chromatogram is less resolved than the AMASE11 chromatogram, while at

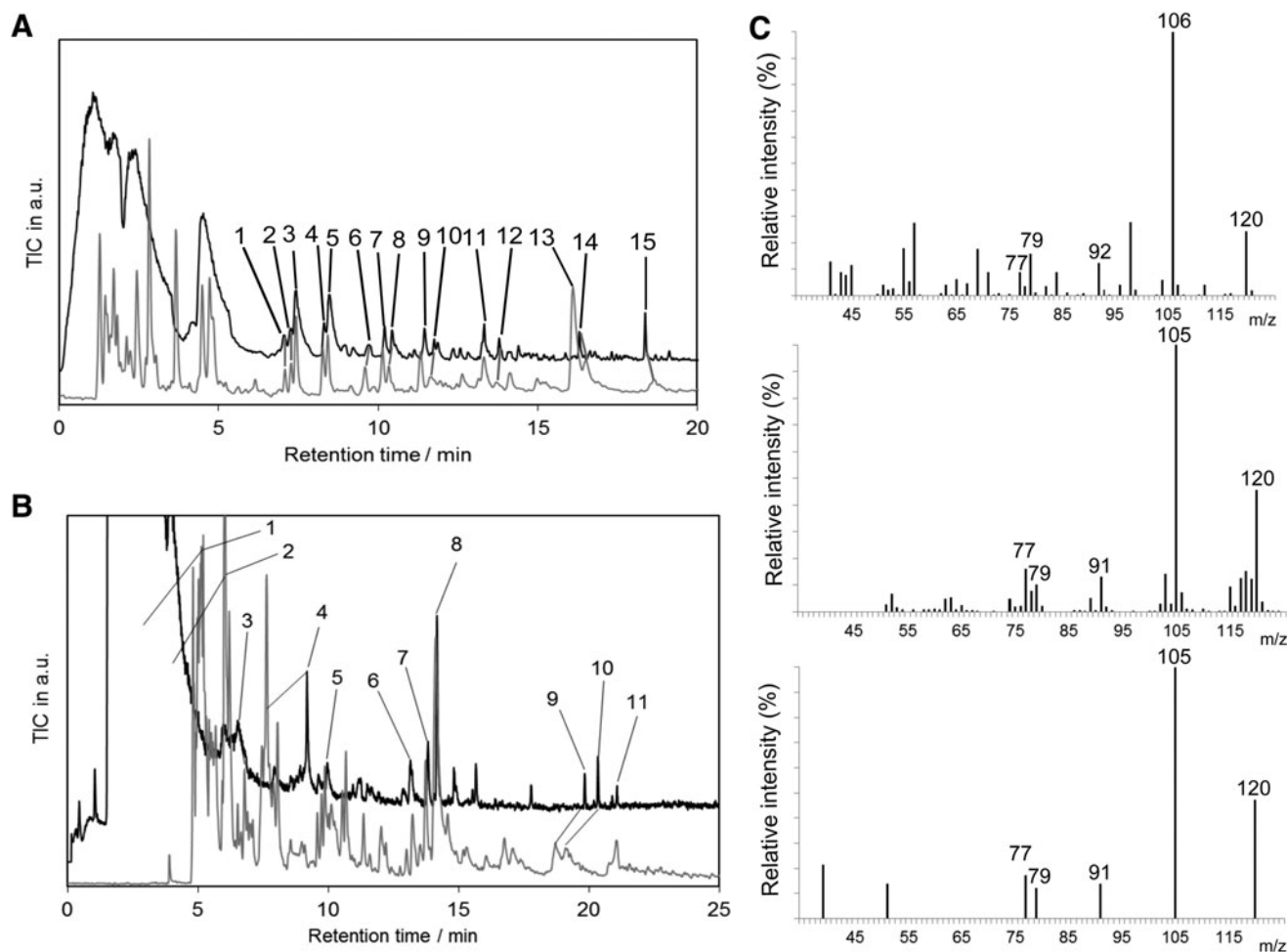


FIG. 4. Comparison of pyrolysis results from the MOMA prototype during AMASE11 and commercial measurements. (A) The TIC of the Butterfinger sample from Colletthøgda with the commercial equipment (top, black) and the MOMA prototype (bottom, gray). The peaks and their respective identification are summarized in Table 2. Only one additional peak (13) is observed in the AMASE chromatogram. This is probably a contamination peak of the Svalbard setup. (B) The TIC of the AM11-162 sample from Botniahalvøya. The black top curve was acquired with the commercial setup, while the gray bottom curve was acquired by using the MOMA prototype. The peaks and their identification are listed in Table 3. (C) Comparison of mass spectra from the Griffin MS (top), the commercial Varian system (middle), and the NIST library (bottom) of the trimethylbenzene (peak 8 in 4A and Table 2) found in the Butterfinger sample. Relative intensity on the y axis is set to the peak with the highest intensity, which is set to 100%. Note that the Griffin MS has an unresolved offset of mass calibration of 1–2 u.

TABLE 2. PEAK ASSIGNMENTS FOR FIGURE 4A

Peak No.	Varian 4000 GCMS		Griffin MS with MOMA GC	
	Main mass peaks (m/z)	Substance	Main mass peaks (m/z)	Substance
1	55, 70, 83, 105, 117	Hydrocarbon	57, 71, 85, 99	Hydrocarbon
2	91, 106	Xylene	92, 107	Xylene
3	91, 106	Xylene	92, 107	Xylene
4	91, 106	Xylene	92, 107	Xylene
5	78, 104	Styrene	79, 105	Styrene
6	281	Column bleeding	281	Column bleeding
7	55, 70, 83, 105, 117	Hydrocarbon	57, 71, 85, 99	Hydrocarbon
8	79, 91, 105, 120	Trimethylbenzene	79, 92, 106, 120	Trimethylbenzene
9	79, 91, 105, 120	Trimethylbenzene	106, 119	Trimethylbenzene
10	79, 91, 103, 117	Alkylbenzene	105, 117	Alkylbenzene
11	57, 70, 83, 105, 117	Hydrocarbon	57, 71, 85	Hydrocarbon
12	105, 115, 116	Indene or ethyltoluene	106, 115	Indene or ethyltoluene
13	N/A	N/A	57, 69, 98, 101, 124	?
14	57, 71, 84	Hydrocarbon	57, 71, 85	Hydrocarbon
15	128	Naphthalene	128	Naphthalene

The TIC of the Butterfinger sample. Note that the Griffin MS has an unresolved offset of mass calibration of 1–2 u.

higher retention times (>5 min) there is virtually a one-to-one correlation between the two systems. This can be explained by the fact that, in the commercial pyrolysis-GCMS, the sample is heated directly in the helium flow of the GC—that is, the pyrolysis and the injection are one single step—while in the prototype (MOMA oven and MOMA GC coupled to Griffin MS), the sample is first heated in a helium flow, and the volatiles are trapped in an injection trap. In a second step, the volatiles are desorbed from the trap and injected into the GCMS. This results in a higher resolution of the MOMA GC than the commercial instrument at low retention times. At higher retention times, the heavier volatiles are better transferred in the commercial GCMS.

The meaning of these results is twofold. First, the measurement obtained on the prototype during AMASE11 can be reproduced in the laboratory. Second, apart from very few exceptions (one prominent peak around 18 min), all peaks found in the chromatograms of the MOMA prototype during AMASE11 are also present in the laboratory data (Figs. 4, 6, and 8), that is, the contamination is the same in both data sets and, since the empty runs with the MOMA prototype during AMASE11 and laboratory equipment are clean, rather small.

TABLE 3. PEAK ASSIGNMENTS FOR FIGURE 4B

Peak No.	Substance
1	Sulfur dioxide
2	Thiophene
3	Methylthiophene
4	Furfural
5	1(2-Furanyl)-ethanone
6	Benzaldehyde
7	Benzofuran
8	Benzonitrile
9	Naphthalene
10	Cyclopenta(c)thiapyran
11	Thienothiophene

The TIC of AM11-162 from Botniahelvøya.

Derivatization. The chromatograms of the derivatized sample of KitKat (Fig. 5), Butterfinger (not shown), and Snickers (not shown) obtained during the AMASE11 campaign on the Griffin GCMS are compared to the chromatograms of these same samples obtained at GSFC on a Thermo Finnigan GCMS (Fig. 7). The on-site results from KitKat, Butterfinger, and Snickers are consistent with the laboratory experiments with various organic species such as carboxylic acids detected above background in KitKat (Fig. 5 and 7), while a lower variety of organics was detected in Snickers and Butterfinger (not shown). This confirms the success of a one-pot, one-step extraction and derivatization of organics from a solid sample, a procedure that will be performed on MOMA. It should be noted that the Griffin GCMS used during AMASE uses an ion trap MS, while the MS used back at GSFC is a quadrupole.

4.3. Laboratory measurements at home institutions

4.3.1. Colletthøgda

Pyrolysis. Snickers, KitKat, and Butterfinger (Fig. 6A, 6B) all contain low amounts of organic material (Butterfinger organic carbon $\leq 0.05\%$ as measured by IR spectroscopy) of mainly aliphatic nature and aromatic nature. Butterfinger contains significantly more organics than the other two samples of both aliphatic and aromatic nature.

The m/z 55, 56, and 57 are diagnostic masses for saturated and unsaturated aliphatic hydrocarbon molecules. The selective ion chromatogram (m/z 55, 56, and 57) of the Butterfinger sample reveals only a single discrete peak, an alkene with a terminal double bond, for each chain length between C8 and C16 (Fig. 6A, 6C). By comparison, in the selective ion chromatogram (m/z 55, 56, 57) of the KitKat (Fig. 6B, 6C) and the Snickers (not shown) samples, there are several peaks for each chain length between C10 and C14. Chain length C10, C12, C13, and C14 dominate, while C11 is present but at a much lower concentration (Fig. 6B, 6C). The chromatogram of Snickers is similar to that of KitKat. The decreased signal intensity at higher chain lengths can be either sample-based (real), procedure-based (cracking of

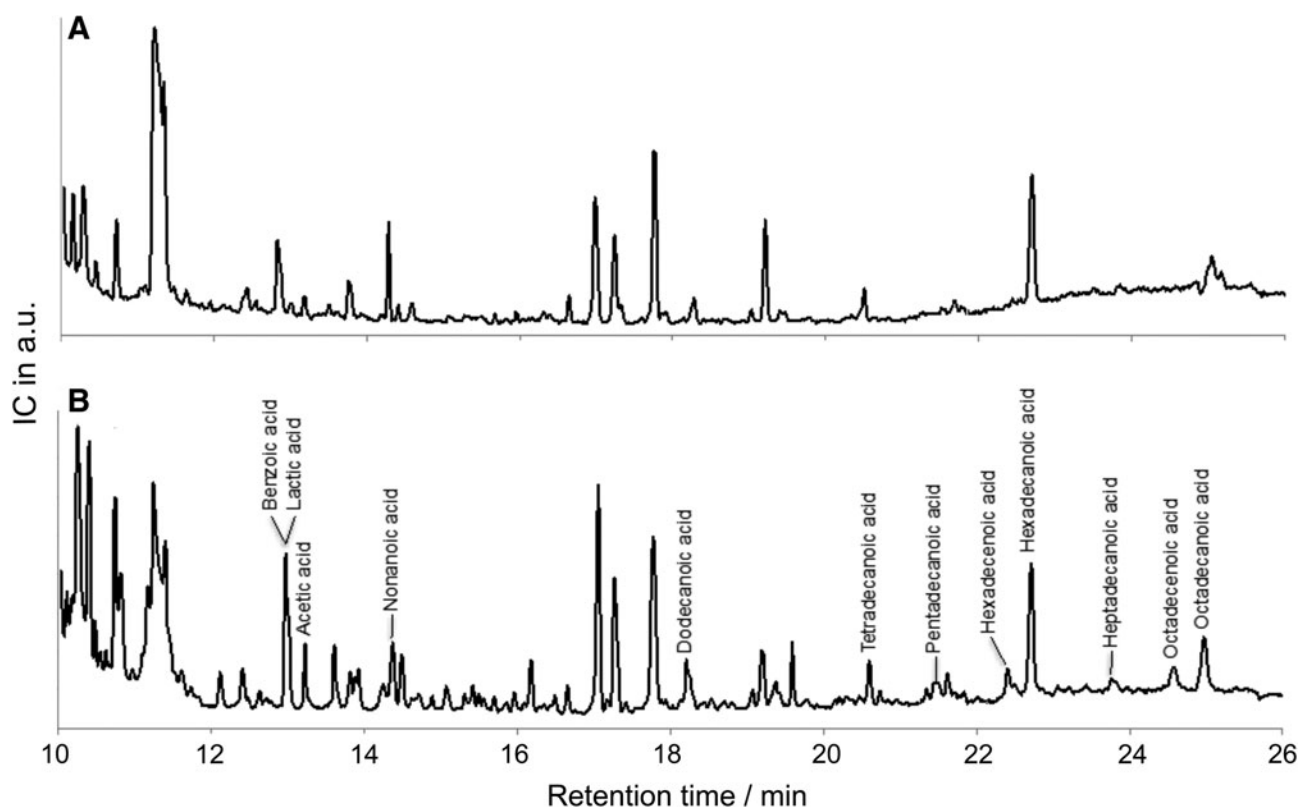


FIG. 5. Derivatization experiments performed during the AMASE11 campaign. (A) Blank and (B) KitKat sample. In the experiments, extraction and derivatization of organics were performed in a single step and single vial as would be in the case of the MOMA experiment.

longer chains by pyrolysis), or instrument-based (longer chains have a lower probability of passing through the GC).

The selective ion chromatograms of the aromatic fractions (m/z 78, 91, 105, 120, 128, and 154) of the Butterfinger and KitKat samples are plotted in Fig. 6D (assignments of peaks in Table 4). If the chromatograms are normalized to the largest peak (benzene, peak 1), it is obvious that Butterfinger contains a higher quantitative abundance of the larger aromatic hydrocarbons than KitKat and Snickers (not shown), though it is of similar qualitative composition. Aromatics found in the samples include benzene, naphthalene, toluene, xylene, and various alkylbenzenes and alkylnaphthalenes. No compounds (aromatic or heterocyclic) containing nitrogen, oxygen, and sulfur were detected in either sample.

Derivatization. In the derivatization experiment, organics were detected in the KitKat (Fig. 7B) and Butterfinger (not shown) samples but not in the Snickers sample (not shown). The organic molecules are detected as their tBDMS ester derivatives but are referred to as their parent molecule before derivatization in the figures and in the text.

The KitKat sample is the most qualitative organic-rich of the three samples and primarily contains a large variety of carboxylic acids, from simple molecules ($<C_{10}$) such as benzoic acid, lactic acid, acetic acid, or nonanoic acid, to heavier long-chain carboxylic acids from dodecanoic (C_{16}) to octadecanoic acids (C_{18}). Interestingly, some of the carboxylic acids display their alkane as well as their alkene form, such as octadecanoic/octadecenoic acid (C_{16}) or hexadecanoic/hexadecenoic acid

(C_{18}). A few dicarboxylic acids were detected, and a significant amount of carboxylic acids could not definitely be identified. In addition to carboxylic acids, a possible long-chain alcohol is present in the sample, extending the family of compounds detected in the samples.

The Butterfinger sample does not contain organic compounds above the background level, apart from some lactic and benzoic acid at low levels of concentration, while the Snickers sample does not contain any organics that could be detected by MTBSTFA derivatization. However, Snickers does contain a significant concentration of sulfur-containing compounds; sulfuric acid is the main compound retrieved after derivatization, and a molecule that is thought to be sulfamate was also detected. The sulfuric acid probably originates from the dissolution of the anhydrite and gypsum of the sample in the water/isopropanol mixture, which is then dried out and derivatized.

Interpretation. Pyrolysis and derivatization show similar results for the samples from Colletthøgda (Figs. 6 and 7B). Both samples, for example, indicate the presence of carboxylic acids in the samples. However, one main difference is that, in the pyrolysis experiment, the most organic-rich sample was Butterfinger, while for the derivatization experiment it was KitKat. A probable explanation for this discrepancy is due to sample heterogeneity. Butterfinger contains veins that could contain a higher concentration of organic material than the cement. Additional pyrolysis experiments of Butterfinger revealed a lower concentration,

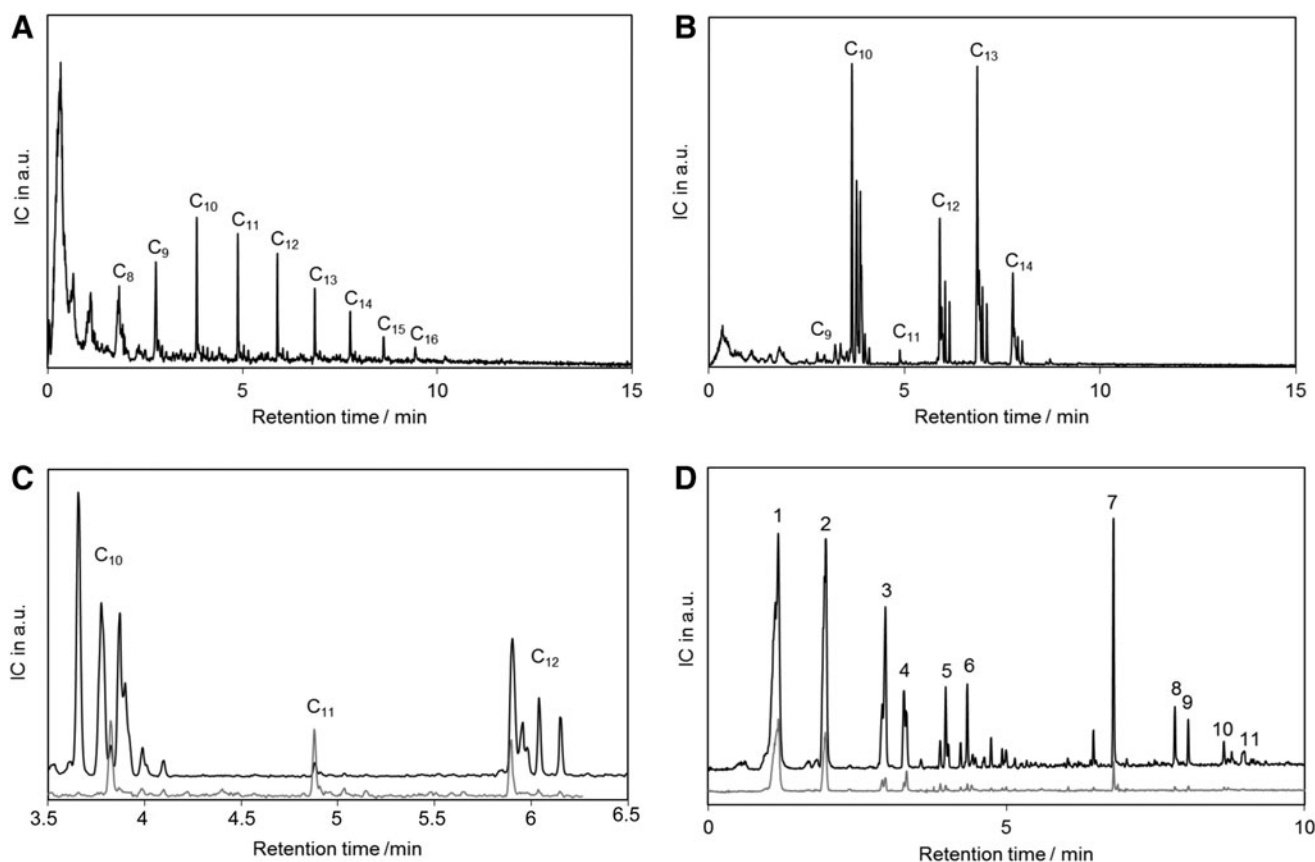


FIG. 6. Pyrolysis (600°C) results from Colletthøgda samples. Selective ion mode (m/z 55, 56, and 57) chromatograms of (A) Butterfinger, (B) KitKat, and (C) Butterfinger (bottom, gray) and KitKat (top, black). (D) Selective ion mode (m/z 78, 91, 105, 120, 128, and 154) of Butterfinger (top, black) and KitKat (bottom, gray). The peaks and their respective identification are listed in Table 4.

but of qualitative similar distribution, of organics (20–50% less) than observed in the initial measurements. Furthermore, the organic composition of Butterfinger and KitKat samples could be different, with more polar compounds present in the KitKat sample and more apolar present in the Butterfinger sample. Derivatization would be more likely to detect polar compounds, while pyrolysis would be more likely to detect apolar compounds.

TABLE 4. PEAK ASSIGNMENTS FOR FIGURE 6D

Peak No.	Substance
1	Benzene
2	Toluene
3	Xylene
4	Ethylbenzene
5	Trimethylbenzene
6	Methylethylbenzene
7	Naphthalene
8	Methylnaphthalene
9	Methylnaphthalene
10	Dimethylnaphthalene
11	Dimethylnaphthalene

Selective ion mode (m/z 78, 91, 105, 120, 128, and 154) chromatogram of Butterfinger and KitKat.

It is likely that some of the molecules found in the pyrolysis experiment are pyrolysis products of compounds detected in the derivatization experiment. For example, the long-chained hydrocarbons detected in the Butterfinger, KitKat, and Snickers samples may have been formed during decarboxylation of carboxylic acids during pyrolysis (Simmonds *et al.*, 1969; Wilson *et al.*, 1983). However, carboxylic acids were only detected in KitKat during the derivatization experiment. Therefore, alternative sources for the long-chained hydrocarbons detected in the pyrolysis experiments must be considered. One possibility would be alkanes that are cleaved during pyrolysis to produce alkenes (Holman *et al.*, 1966; Henderson *et al.*, 1968). A further possibility could be fossil material. Chromatograms obtained from pyrolysis experiments of fossils often contain a significant amount of alkanes/alkenes (Gupta *et al.*, 2007). The alkanes/alkenes are thought to originate from cross-linking of these types of molecules into the original organic material in the fossil (Gupta *et al.*, 2007).

In the derivatization experiment, saturated carboxylic acids (C14–C18) were found in KitKat. The two most abundant carboxylic acids, C16 and C18, were also found in their unsaturated form, such as octadecenoic acid (C16) and hexadecenoic acid (C18). Saturated and unsaturated fatty acids, especially palmitoleic (C16) and oleic (C18) acids, are important components of the cell membrane of bacteria and eukaryotes (Peters *et al.*, 2005).

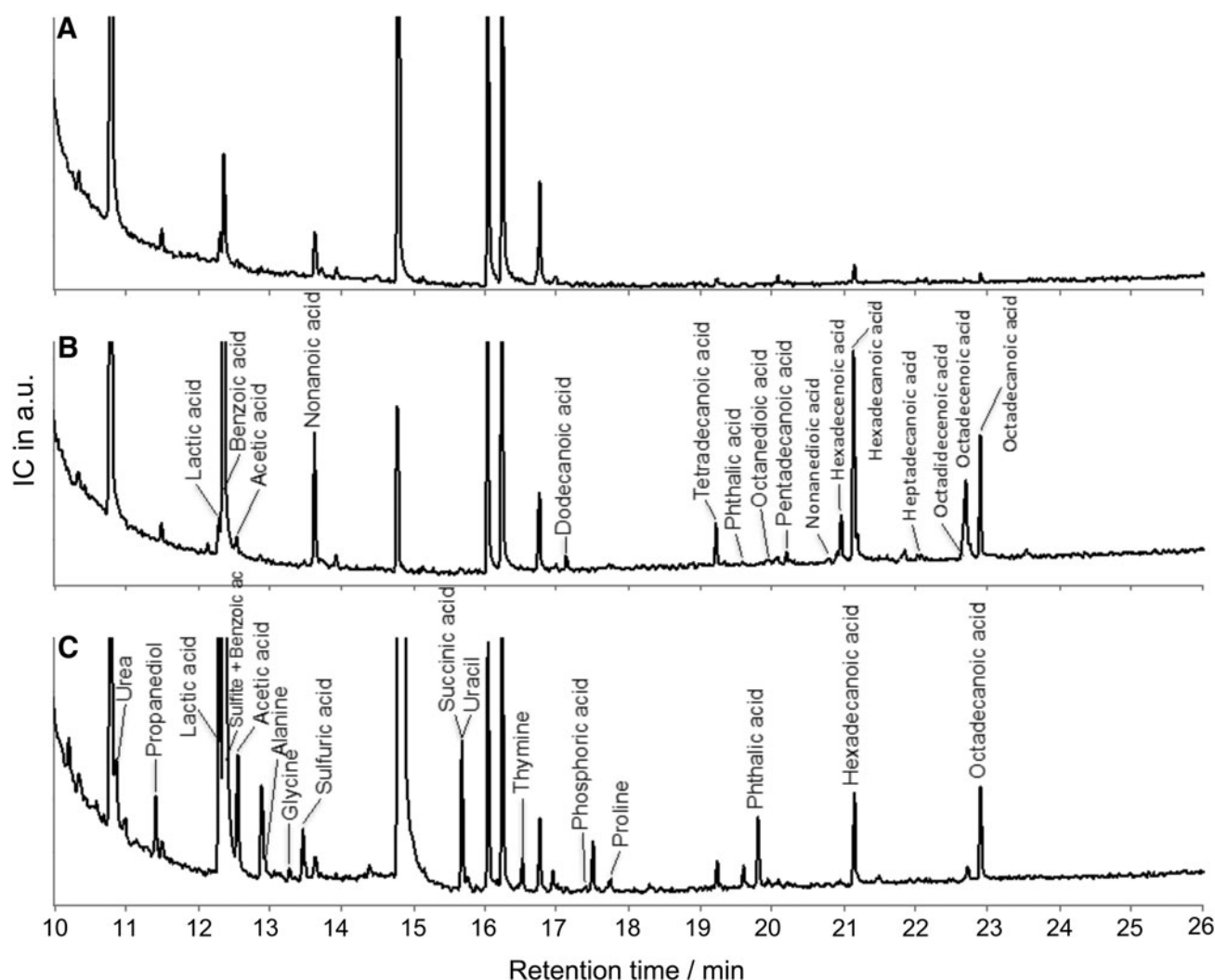


FIG. 7. Derivatization experiments of (A) blank, (B) KitKat, and (C) Botniahavøya samples. The chromatogram displays derivatized molecules; however, the names shown on the chromatogram are the parent molecules present in the sample before derivatization. The spectra is a post-treated selective ion mode that includes m/z 105, 171, 179, 189, 198, 215, 218, 232, 253, 257, 263, 283, 285, 289, 297, 299, 311, 313, 325, 327, 337, 339, 341, 359, 383, and 414 to enhance the signal-to-noise ratio for the molecules of interest described here.

In addition to the carboxylic acids, benzoic and phthalic acids were detected in significant concentration in the derivatization of KitKat and at lower abundances in Butterfinger. Those compounds are typical stable degradation products of organic molecules due to oxidation in the soil (Benner *et al.*, 2000; Navarro-González *et al.*, 2006). In addition to being an oxidation product, phthalic acids may also originate from phthalates, a common organic contaminant (Peters *et al.*, 2005). Acetic acid, a simple carboxylic acid, is also present above background level in the derivatization experiment of KitKat. This molecule is believed to be a metastable decomposition product of various components such as amines and primary alcohols (Benner *et al.*, 2000). It is possible that some of the aromatic compounds found in the pyrolysis experiment, such as benzene, are pyrolysis products of the benzoic and phthalic acids found in the derivatization experiment.

The detection of the various organic molecules in KitKat, Butterfinger, and Snickers—including the extended range of long-chained hydrocarbons in pyrolysis and carboxylic acids

in derivatization—and the absence of amino acids might indicate the presence of a once-thriving biota in the rocks at Colletthøgda. The presence of carboxylic acids in KitKat indicates either excellent preservation potential of now nonviable organisms or the presence of an active biota. Fatty acids have a better survival potential in sediments and fossils than many other biomolecules (Lee *et al.*, 2004; Petsch, 2011). For example, fatty acids have been detected in the Messel oil shale from the Eocene and in fossils from the Miocene (Lockheart *et al.*, 2000; Thiel and Sjövall, 2011). They have even been detected in carbonate samples from the Devonian (Melendez *et al.*, 2013). However, preservation of fatty acids usually occurs under very specific conditions such as euxinia, that is, anoxic conditions in the presence of H_2S . It is possible that the carboxylic acids found in samples from Colletthøgda have been preserved since the Permian, though it is more likely they are remains of more recent colonization. Laboratory experiments have also shown that many types of organic molecules, including

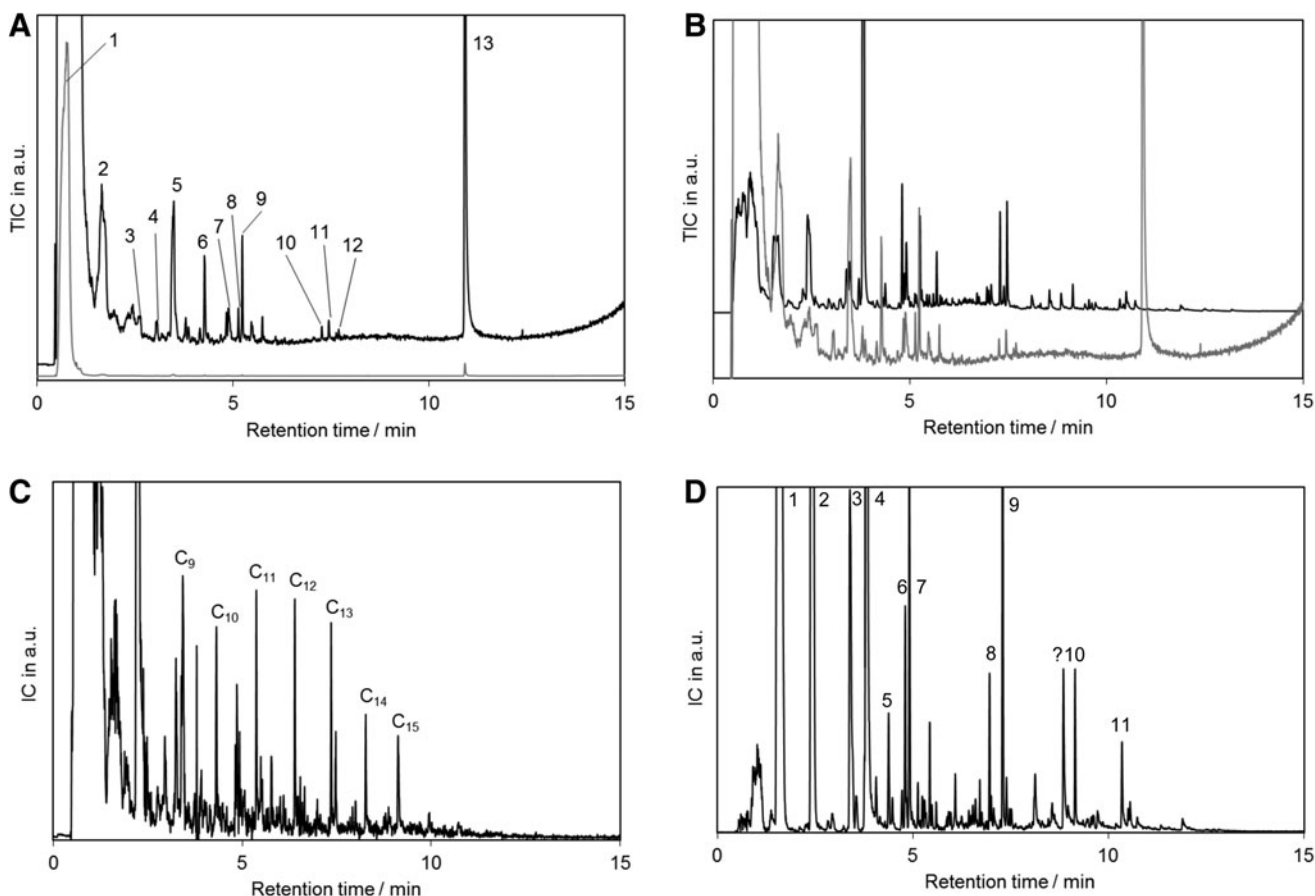


FIG. 8. Pyrolysis results from Botniahalvøya samples. (A) The TIC of the Botniahalvøya bulk sample pyrolyzed at a temperature of 900°C. Bottom gray chromatogram represents blank, and top black represents sample. The peaks and their respective identification are listed in Table 5. (B) Comparison of the TICs of the Botniahalvøya bulk sample (bottom, gray) and the scraped-off outer layer of the sample (top, black). Selective ion mode of Botniahalvøya outer layer (C) (m/z 55, 56, and 57) and (D) (m/z 78, 91, 105, 120, 128, and 154). The peaks and their respective identification are listed in Table 6.

long-chained hydrocarbons and carboxylic acids, may also form abiotically under hydrothermal conditions (McCormell *et al.*, 1999; Rushdi and Simoneit, 2001). However, the mineralogy of the samples of Colletthøgda and the overall geology of the site as a shallow marine environment suggest a low-temperature environment, and the dominance of C16 and C18 carboxylic acids favors a biological origin for these molecules.

4.3.2. Botniahalvøya

Pyrolysis. In the pyrolysis experiment, the TIC of the bulk sample is dominated by sulfur-containing compounds such as S₆, SO₂, and benzothiophenes (Fig. 8A, Table 5). It is not clear whether those compounds were created during pyrolysis (Simmonds, 1970) or are breakdown products of larger molecules present in the sample. In any case, the release of SO₂ and S₆ requires a significant sulfur source (Faber, 1920; Beerwald, 1950). To investigate whether such a source of sulfur exists in the Botniahalvøya sample, SEM/EDX analysis was performed (Fig. 9). Analysis results show strong peaks from iron, silica, and oxygen with smaller peaks of calcium, titanium, magnesium, and aluminum present, especially in the outer layer. Sulfur is present in localized spots (not shown). Sulfide minerals (sulfates can-

not produce S₆) present might produce enough sulfur to generate various reactions during the pyrolysis and generate the sulfur-containing organic molecules (McAdam *et al.*, 2014). When the visible black layer—the proposed desert varnish—was scraped from the rock and analyzed

TABLE 5. PEAK ASSIGNMENTS FOR FIGURE 8A

Peak No.	Substance
1	Sulfur dioxide
2	Thiophene
3	Methylthiophene
4	2-Furancarboxitrile
5	Furfural
6	1-(2-Furanyl)-ethanone
7	Benzaldehyde
8	Benzofuran
9	Benzonitrile
10	Naphthalene
11	Cyclopenta(c)thiapyran
12	Thieno(2,3-b)thiophene
13	S ₆

The TIC of AM11-162 from Botniahalvøya.

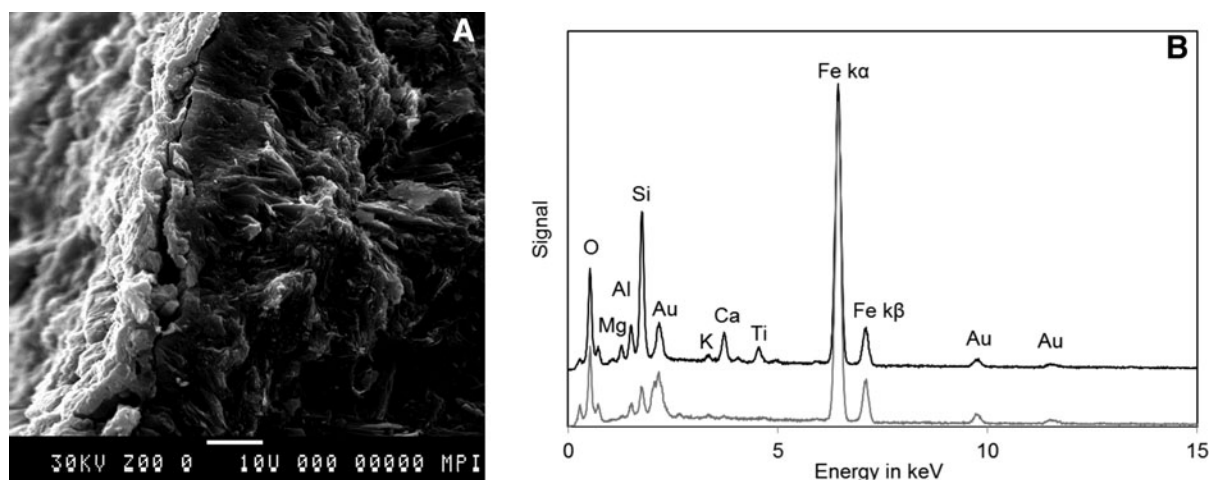


FIG. 9. (A) Secondary electron image of a freshly broken piece of AM11-162 from Botniahalvøya. The outside crust of the rock is seen left in the image. (B) EDX spectra of the inside of AM11-162 (black, top) and the outside crust (gray, bottom). The gold signal is due to the coating necessary to produce good SEM images. Quantitative analysis is difficult, but a clear trend is visible. The inside of the rock contains Fe, Ti, Ca, Si, Al, and O. The outside is reduced in most constituents apart from Fe and O.

separately, the TIC (Fig. 8B) shows much higher levels of organic material with virtually no sulfur in the form of S_6 and SO_2 . The dominant molecule in the chromatogram is styrene. This suggests that most of the organics present in the sample are located in the outer layer.

The selective ion chromatogram (m/z 55, 56, 57) of the outer layer of the sample shows a range of alkenes with terminal double bonds of chain length between C9 and C17 (Fig. 8C).

The selective ion chromatogram (m/z 78, 91, 105, 120, 128, and 154) of the outer layer of the Botniahalvøya sample displays sulfur-containing compounds such as thiophenes and benzothiophenes (Fig. 8D), in addition to the normal pyrolysis products of organic material such as benzene, toluene, and naphthalene. Other detected compounds in the sample included nitrogen- and oxygen-containing molecules such as nitriles and furans, which is characteristic for materials that contain living biota as opposed to more ancient material that primarily contains hydrocarbons and aromatics (Simmonds, 1970). Molecules detected in Botniahalvøya include furfuraldehyde, which may be created during pyrolysis of sugars such as glucose and ribose (Kato, 1967; Simmonds, 1970; Wilson *et al.*, 1983). Other furan derivatives such as benzofuran and 1-(2-furanyl)-ethanone were also detected (Fig. 8D, Table 6) (Simmonds, 1970). In addition, nitrogen-containing compounds such as benzonitrile and 2-furancarbonitrile were detected. Benzonitrile is one of the molecules created during the pyrolysis of phenylalanine (Chiavari and Galletti, 1992) and has been found to be indicative of this type of amino acid (Gutteridge and Norris, 1979).

Derivatization. The derivatization experiment of Botniahalvøya after water/isopropanol extraction is presented in Fig. 7C. As in the pyrolysis experiment, this sample displayed the richest diversity of both organic and inorganic molecules. Detected inorganic molecules include S-containing sulfite and sulfuric acid, and detected organic molecules include N- and/or O-containing molecules such as urea, carboxylic acids, nucleobases, and amino acids.

Amino acids such as glycine, alanine, and proline are identified with high confidence, and the presence of valine is also suspected. These amino acids would not be seen in the pyrolysis experiment, as they mostly decompose to smaller molecules such as CO_2 , NH_3 and small nitriles, imines, and amines (Simmonds *et al.*, 1972; Chiavari and Galletti, 1992), which was not resolved by the GC conditions used in this experiment. The tentative detection of phenylalanine in the pyrolysis experiment and its absence in the derivatization experiment may indicate a low efficiency of the extraction procedure used here for this type of molecule. Indeed, the procedure used for extraction (water/isopropanol) was chosen because it is a nonselective method that allows a large number of compound families to be extracted. However, the specific efficiency toward different amino acids is not known.

Thymine and uracil, two nucleobases of DNA and RNA, were detected in the sample in the derivatization experiment. These two bases are pyrimidine bases, and their detection rather than that of the purine bases was expected due to the smaller size of the molecule. Nucleobases were not

TABLE 6. PEAK ASSIGNMENTS FOR FIGURE 8D

Peak No.	Substance
1	Benzene
2	Toluene
3	Ethylbenzene
4	Styrene
5	Benzene 2-propenyl
6	Methylstyrene
7	Benzaldehyde
8	Azulene
9	Naphthalene
10	Biphenyl
11	Bibenzyl

Selective ion chromatogram (m/z 78, 91, 120, 128, 154) of AM11-162 from Botniahalvøya.

detected in the pyrolysis experiment, as they decompose to smaller molecules such as HCN (Gutteridge and Norris, 1979), which was not resolved in the pyrolysis experiment.

Moreover, phosphoric acid was detected. Phosphoric acid is a key component of the DNA skeleton, and its presence, along with nucleobases and amino acids, is a definite sign of a living biota, as the preservation potential of DNA is poor in most sediments.

Interpretation. The Botniahalvøya sample is rich in organics, especially in the outer layer. The detected organic molecules include sugars, amino acids, carboxylic acids, and nucleobases as indicated by both the pyrolysis and derivatization experiments. The detection of these compounds indicates the presence of extant biota or a recently extinct one, as many of these compounds usually start degrading directly after the death of organisms (Brocks and Grice, 2011; Petsch, 2011). Visual inspection and XRD analysis of the sample indicated the presence of clays and iron oxides, with manganese oxides on the samples indicating the presence of desert varnish. Living organisms are often associated with desert varnish.

The Botniahalvøya sample is heterogenic and probably contains sulfide minerals based on the detection of S_6 in the pyrolysis experiment and its oxidized form, sulfite, in derivatization. Evolved gas analysis data obtained on the same sample during the AMASE campaign indicate the possible presence of pyrite in the sample (Amy McAdam, personal communication). Also, results of the EDX analysis indicate the presence of localized sulfur, possibly in the form of pyrite in this sample.

5. Discussion

5.1. Implications for exploration of the organic content of Mars

Laboratory instrument test stands have been developed to allow modeling of various MOMA subsystems, but at the time of deployment no overall breadboard was currently available. Before AMASE11, the MOMA GC had only been tested on a stand-alone base (Goetz *et al.*, 2011b) and had never been coupled to a MS. The coupling of the MOMA GC with a commercial MS during AMASE11 was the first successful interfacing of the MOMA GC with a MS.

The results obtained from samples from Svalbard indicate that the current setup of MOMA for pyrolysis and the MTBSTFA derivatization experiment allows for the detection of different types of inorganic and organic molecules in an extended range of Mars analog samples. The accordance of the results obtained in the laboratory and during the expedition is very promising regarding both pyrolysis and derivatization (Mahaffy *et al.*, 2012). The derivatization experiments also show that similar results are obtained when the extraction and derivatization are performed in successive steps or in a single step. Together with information about the mineralogy of the samples and overall geology of the sample site, it is also possible to draw some conclusions on the source of organics (abiotic or biotic origin, extinct or extant life) and how these organics might have been affected by later geological events after the deposition of the organics. Another valuable lesson from these experiments is the complementary data that can be obtained by combining these two modes, which will reveal important information

on any complex organic molecules present in future samples analyzed by MOMA.

5.2. Future work

The tested MOMA prototype is only a subset of the final instrument, which will contain three additional GC columns, additional derivatization experiments, and a laser desorption ionization mass spectrometer (LDI-MS) (Goetz *et al.*, 2011a). The MS used in this study was a commercial ion trap and is different from the ion trap being built for MOMA. The additional columns combined with the range of possible derivatization experiments are expected to yield crucial information about more polar organic components that are important for life such as amino acids, nucleobases, and sugars. In addition, measurements of the chirality of these molecules might be determined by using the enantioselective derivatization reagent (DMF-DMA) and the chiral column on the flight instrument, which were unfortunately not available for this field test. Furthermore, the LDI-MS will have the capacity to detect larger molecules and provide information on the mineralogy as a complement to the Raman of ExoMars. All these different components of the MOMA instrument need to be tested prior to flight.

Depending on the landing site, the joint ESA/Roscosmos ExoMars mission could analyze a wide variety of soils. These soils could include basalt, carbonate, sulfate, phyllosilicate, and so on and could contain a wide range of hydrated or oxidative species such as hydrogen peroxides and perchlorates as detected by Phoenix (Hecht *et al.*, 2009). Therefore, it is very important to understand how the pyrolysis, derivatization reactions, and LDI-MS behave in the presence of these species, which is why multiple Mars analog samples must be analyzed *a priori* to understand the behavior of the organics under the conditions of the experiments. Future work will combine and test all the aspects of MOMA in a Mars-relevant terrestrial environment. Although valuable information can be obtained from the study of mineral mixtures spiked with organics, they cannot completely replace natural samples that are tested with a scientific hypothesis in mind. It is therefore of great importance to study samples from Mars analog environments on Earth.

6. Conclusion

The first successful coupling of the MOMA GC prototype with a commercial MS was accomplished during the AMASE11 field campaign. Pyrolysis and derivatization experiments were performed under MOMA-like conditions on Mars analog samples collected in Svalbard. These experiments have been replicated in the laboratory with the use of commercial instruments.

The diversity and repartition of organic molecules detected at Colletthøgda, including aromatics, long-chained hydrocarbons, and carboxylic acids, suggest the presence of once-living biota in these samples. The organic compounds detected at Botniahalvøya include carboxylic acids and nitrogen- and oxygen-containing compounds such as amino acids, nucleobases, and sugars. The presence of these compounds indicates the presence of living biota in the sample. A significant source of sulfur is also present at Botniahalvøya.

These results obtained in the field and in the laboratory indicate that, together, the MOMA pyrolysis and derivatization experiments will aid researchers in their efforts to answer questions regarding the organic chemistry of Mars' samples. This is extremely positive and promising for the ExoMars mission goal of searching for biosignatures.

Appendix: The AMASE11 Team

Eivind Aksnes, Earth and Planetary Exploration Services.
Hans E.F. Amundsen, Earth and Planetary Exploration Services.

Arnold Bauer, Joanneum Research.

Dave Blake, NASA Ames Research Center.

Tom Bristow, NASA Ames Research Center.

Caroline Freissinet, NASA Goddard Space Flight Center.

Mihaela Glamoclija, Carnegie Institution of Washington.

Fred Goesmann, Max Planck Institute for Solar System Research.

Garret Huntress, Carnegie Institution of Washington.

Fabien Jordan, SEI-Swiss Engineering Institute.

Laureline Josset, SEI-Swiss Engineering Institute.

Ingrid Kårstad, Earth and Planetary Exploration Services.

Amy McAdam, NASA Goddard Space Flight Center.

Francis McCubbin, University of New Mexico.

Kjell Ove Storvik, Earth and Planetary Exploration Services.

Stephen Pugh, Aberysthwyth University.

Nicole Schmitz, DLR German Aerospace Center.

Sandra Siljeström, SP Technical Research Institute of Sweden.

Pablo Sobron, Washington University in Saint Louis.

Steve Squyres, Cornell University.

Verena Starke, Carnegie Institution of Washington.

Andrew Steele, Carnegie Institution of Washington.

Jen Stern, NASA Goddard Space Flight Center.

Tor Viscor, Kayser-Threde GmbH.

Acknowledgments

The authors would like to thank David Blake and Tom Bristow at NASA Ames Research Center and Steve Squyres at Cornell University for the collection and interpretation of the XRD/XRF data. The authors also would like to thank Amy McAdam at GSFC for her AMASE11 evolved gas analysis work and insights on the sulfur story, the CLUPI team for providing images of the samples, and the two anonymous reviewers for their useful comments that improved the initial manuscript.

The financial support of AMASE by ESA is gratefully acknowledged. The participation of F.G. was supported by Deutsches Zentrum für Luft- und Raumfahrt (DLR; FKZ: 50QX1001). The participation of S.S. in AMASE11 was supported by MPS and a Swedish National Space Board travel grant. S.S. was further supported by the Deep Carbon Observatory (A. Steele, PI), the postdoctoral fellowship program of the Geophysical Laboratory, Carnegie Institution of Washington, and the Swedish National Space Board (Contract No. 121/11). The participation of C.F. was supported by GSFC. AMASE11 was strongly supported by NASA ASTEP (A. Steele, PI).

Author Disclosure Statement

No competing financial interests exist.

Abbreviations

AMASE, Arctic Mars Analog Svalbard Expedition.

DMF, dimethylformamide.

DMF-DMA, dimethylformamide dimethylacetal.

EDX, energy-dispersive X-ray.

GC, gas chromatograph.

GSFC, Goddard Space Flight Center.

LDI-MS, laser desorption ionization mass spectrometer.

MOMA, Mars Organic Molecule Analyzer.

MPS, Max Planck Institute for Solar System Research.

MS, mass spectrometer.

MTBSTFA, *N,N*-methyl-*tert*-butyl-dimethylsilyltrifluoroacetamide.

R/V, research vessel.

SAM, Sample Analysis at Mars.

SEM, scanning electron microscope.

tBDMS, *tert*-butyldimethylsilyl.

TIC, total ion chromatogram.

XRD, X-ray diffraction.

XRF, X-ray fluorescence.

References

- Arvidson, R.E., Gooding, J.L., and Moore, H.J. (1989) The martian surface as imaged, sampled, and analyzed by the Viking landers. *Rev Geophys* 27:39–60.
- Aubrey, A., Cleaves, H.J., Chalmers, J.H., Skelley, A.M., Mathies, R.A., Grunthaner, F.J., Ehrenfreund, P., and Bada, J.L. (2006) Sulfate minerals and organic compounds on Mars. *Geology* 34:357–360.
- Baglioni, P., Fisackerly, R., Gardini, B., Gianfiglio, G., Pradier, A.L., Santovincenzo, A., Vago, J.L., and van Winnendael, M. (2006) The Mars exploration plans of ESA. *IEEE Robot Autom Mag* 13:83–89.
- Beerwald, A. (1950) Über die Reduktion von Magnesiumsulfat mit Kohle. *Zeitschrift für anorganische Chemie* 261:52–62.
- Benner, S.A., Devine, K.G., Matveeva, L.N., and Powell, D.H. (2000) The missing organic molecules on Mars. *Proc Natl Acad Sci USA* 97:2425–2430.
- Bibring, J.-P., Langevin, Y., Mustard, J.F., Poulet, F., Arvidson, R., Gendrin, A., Gondet, B., Mangold, N., Pinet, P., Forget, F., Berthé, M., Gomez, C., Jouglet, D., Soufflot, A., Vincendon, M., Combes, M., Drossart, P., Encrenaz, T., Fouchet, T., Mercurio, R., Belluci, G., Altieri, F., Formisano, V., Capaccioni, F., Cerroni, P., Coradini, A., Fonti, S., Korabiev, O., Kottsov, V., Ignatiev, N., Moroz, V., Titov, D., Zasova, L., Loiseau, D., Pinet, P., Douté, S., Schmitt, B., Sotin, C., Hauber, E., Hoffmann, H., Jaumann, R., Keller, U., Arvidson, R., Duxbury, T., Forget, F., and Neukum, G. (2006) Global mineralogical and aqueous Mars history derived from OMEGA/Mars Express data. *Science* 312:400–404.
- Biemann, K. (2007) On the ability of the Viking gas chromatograph-mass spectrometer to detect organic matter. *Proc Natl Acad Sci USA* 104:10310–10313.
- Biemann, K. and Bada, J.L. (2011) Comment on “Reanalysis of the Viking results suggests perchlorate and organics at mid-latitudes on Mars” by Rafael Navarro-González et al. *J Geophys Res Planets* 116, doi:10.1029/2011je003869.
- Biemann, K., Oro, J., Toulmin, P., III, Orgel, L.E., Nier, A.O., Anderson, D.M., Simmonds, P.G., Flory, D., Diaz, A.V., Rushneck, D.R., Biller, J.E., and Lafleur, A.L. (1977) The search for organic substances and inorganic volatile compounds in the surface of Mars. *J Geophys Res* 82:4641–4658.

- Bishop, J.L., Franz, H.B., Goetz, W., Blake, D.F., Freissinet, C., Steininger, H., Goesmann, F., Brinckerhoff, W.B., Getty, S., Pinnick, V.T., Mahaffy, P.R., and Dyar, M.D. (2013) Co-ordinated analyses of Antarctic sediments as Mars analog materials using reflectance spectroscopy and current flight-like instruments for CheMin, SAM and MOMA. *Icarus* 224:309–325.
- Blake, D. (2010, September) A historical perspective of the development of the CheMin instrument for the Mars Science Laboratory (MSL '11) mission. *Geochemical News* 144. Available online at <http://www.geochemsoc.org/publications/geochemicalnews/gn144sep10/ahistoricalperspectiveofth>.
- Brooks, J.J. and Grice, K. (2011) Biomarkers (molecular fossils). In *Encyclopedia of Geobiology*, Vol. 1, edited by J. Reitner and V. Thiel, Springer, Dordrecht, the Netherlands, pp 147–167.
- Chiavari, G. and Galletti, G.C. (1992) Pyrolysis–gas chromatography/mass spectrometry of amino acids. *J Anal Appl Pyrolysis* 24:123–137.
- Dallmann, W.K., editor. (1999) *Lithographic Lexicon of Svalbard*, Norsk Polarintitutt, Tromsø, Norway.
- Dallmann, W.K., editor. (2002) *Bedrock Map of Svalbard and Jan Mayen*, Norsk polarinstitutt, Örebro, Norway.
- Dartnell, L.R., Desorgher, L., Ward, J.M., and Coates, A.J. (2007) Martian sub-surface ionising radiation: biosignatures and geology. *Biogeosci Discuss* 4:455–492.
- DiGregorio, B.E. (2002) Rock varnish as a habitat for extant life on Mars. In *Proc SPIE* 4495, doi:10.1117/12.454750.
- Dorn, R.I., editor. (1998) *Rock Coatings*, Elsevier, Amsterdam.
- Dorn, R.I. and Oberlander, T.M. (1981) Microbial origin of desert varnish. *Science* 213:1245–1247.
- Ehlmann, B.L., Mustard, J.F., Murchie, S.L., Poulet, F., Bishop, J.L., Brown, A.J., Calvin, W.M., Clark, R.N., Des Marais, D.J., Milliken, R.E., Roach, L.H., Roush, T.L., Swayze, G.A., and Wray, J.J. (2008) Orbital identification of carbonate-bearing rocks on Mars. *Science* 322:1828–1832.
- Eigenbrode, J., Benning, L.G., Maule, J., Wainwright, N., Steele, A., and Amundsen, H.E. (2009) A field-based cleaning protocol for sampling devices used in life-detection studies. *Astrobiology* 9:455–465.
- Faber, A. (1920) Die Reduktion von Kieserit durch Kohle. *Journal für Praktische Chemie* 100:116–135.
- Flynn, G.J. (1996) The delivery of organic matter from asteroids and comets to the early surface of Mars. *Earth Moon Planets* 72:469–474.
- Freissinet, C., Glavin, D.P., Miller, K.E., Buch, A., Brunner, A.E., Cabane, M., Eigenbrode, J.L., Fairen, A.G., Kashyap, S., Martin, M., Navarro-González, R., Steele, A., Summons, R.E., Szopa, C., Mahaffy, P.R., and the SAM and MSL science teams. (2014) From SAM instrument background to martian signal: challenges of solid sample analyses on Mars [abstract 2796]. In *45th Lunar and Planetary Science Conference Abstracts*, Lunar and Planetary Institute, Houston.
- Gendrin, A., Mangold, N., Bibring, J.-P., Langevin, Y., Gondet, B., Poulet, F., Bonello, G., Quantin, C., Mustard, J., Arvidson, R., and LeMouélis, S. (2005) Sulfates in martian layered terrains: the OMEGA/Mars Express view. *Science* 307:1587–1591.
- Glavin, D.P., and Bada, J.L. (1998) Isolation of amino acids from natural samples using sublimation. *Anal Chem* 70:3119–3122.
- Glavin, D.P., Freissinet, C., Miller, K.E., Eigenbrode, J.L., Brunner, A.E., Buch, A., Sutter, B., Archer, P.D., Atreya, S.K., Brinckerhoff, W.B., Cabane, M., Coll, P., Conrad, P.G., Coscia, D., Dworkin, J.P., Franz, H.B., Grotzinger, J.P., Leshin, L.A., Martin, M.G., McKay, C., Ming, D.W., Navarro-González, R., Pavlov, A., Steele, A., Summons, R.E., Szopa, C., Teinturier, S., and Mahaffy, P.R. (2013) Evidence for perchlorates and the origin of chlorinated hydrocarbons detected by SAM at the Rocknest aeolian deposit in Gale Crater. *J Geophys Res Planets* 118:1955–1973.
- Goetz, W., Steininger, H., and Goesmann, F. (2011a) Searching for martian organics with the Mars Organic Molecule Analyzer (MOMA) aboard ExoMars-2018. In *EPSC-DPS Joint Meeting 2011 Abstracts*, Europlanet, Paris, p 1281.
- Goetz, W., Steininger, H., Steinmetz, E., Bierwirth, M., Goesmann, F., Philippon, C., Lustreant, B., Szopa, C., Buch, A., Amundsen, H., Fogel, M., and Steele, A. (2011b) Mars Organic Molecule Analyzer (MOMA) field test as part of the AMASE 2010 Svalbard expedition [abstract 2460]. In *42nd Lunar and Planetary Science Conference Abstracts*, Lunar and Planetary Institute, Houston.
- Gooding, J.L. (1992) Soil mineralogy and chemistry on Mars: possible clues from salts and clays in SNC meteorites. *Icarus* 99:28–41.
- Guinness, E.A., Arvidson, R.E., Clark, I.H., and Shepard, M.K. (1997) Optical scattering properties of terrestrial varnished basalts compared with rocks and soils at the Viking lander sites. *J Geophys Res Planets* 102:28687–28703.
- Gupta, N.S., Briggs, D.E., Collinson, M.E., Evershed, R.P., Michels, R., Jack, K.S., and Pancost, R.D. (2007) Evidence for the *in situ* polymerisation of labile aliphatic organic compounds during the preservation of fossil leaves: implications for organic matter preservation. *Org Geochem* 38:499–522.
- Gutteridge, C.S. and Norris, J.R. (1979) The application of pyrolysis techniques to the identification of micro-organisms. *J Appl Bacteriol* 47:5–43.
- Hecht, M.H., Kounaves, S.P., Quinn, R.C., West, S.J., Young, S.M.M., Ming, D.W., Catling, D.C., Clark, B.C., Boynton, W.V., Hoffman, J., DeFlores, L.P., Gospodinova, K., Kapit, J., and Smith, P.H. (2009) Detection of perchlorate and the soluble chemistry of martian soil at the Phoenix lander site. *Science* 325:64–67.
- Henderson, W., Eglinton, G., Simmonds, P., and Lovelock, J.E. (1968) Thermal alteration as a contributory process to the genesis of petroleum. *Nature* 219:1012–1016.
- Holman, R., Deubig, M., and Hayes, H. (1966) Pyrolysis chromatography of lipids. I. Mass spectrometric identification of pyrolysis products of hydrocarbons. *Lipids* 1:247–253.
- Jull, A.J.T., Courtney, C., Jeffrey, D.A., and Beck, J.W. (1998) Isotopic evidence for a terrestrial source of organic compounds found in martian meteorites Allan Hills 84001 and Elephant Moraine 79001. *Science* 279:366–369.
- Kato, K. (1967) Pyrolysis of cellulose part III. Comparative studies of the volatile compounds from pyrolysates of cellulose and its related compounds. *Agr Biol Chem* 31:657–663.
- Klein, H.P. (1999) Did Viking discover life on Mars? *Orig Life Evol Biosph* 29:625–631.
- Knapp, D.R., editor. (1979) *Handbook of Analytical Derivatization Reactions*, John Wiley & Sons, New York.
- Kounaves, S.P., Carrier, B.L., O'Neil, G.D., Stroble, S.T., and Claire, M.W. (2014) Evidence of martian perchlorate, chlorate, and nitrate in Mars meteorite EETA79001: implications for oxidants and organics. *Icarus* 229:206–213.
- Langevin, Y., Poulet, F., Bibring, J.-P., and Gondet, B. (2005) Sulfates in the north polar region of Mars detected by OMEGA/Mars Express. *Science* 307:1584–1586.

- Lanza, N., Ollila, A., Cousin, A., Hardgrove, C., Wiens, R., Mangold, N., Nachon, M., Fabre, C., Bridges, N., and Johnson, J. (2014) Manganese trends with depth on rock surfaces in Gale Crater, Mars [abstract 2599]. In *45th Lunar and Planetary Science Conference Abstracts*, Lunar and Planetary Institute, Houston.
- Lee, C., Wakeham, S., and Arnosti, C. (2004) Particulate organic matter in the sea: the composition conundrum. *AMBIO: A Journal of the Human Environment* 33:565–575.
- Leshin, L.A., Mahaffy, P.R., Webster, C.R., Cabane, M., Coll, P., Conrad, P.G., Archer, P.D., Atreya, S.K., Brunner, A.E., Buch, A., Eigenbrode, J.L., Flesch, G.J., Franz, H.B., Freissinet, C., Glavin, D.P., McAdam, A.C., Miller, K.E., Ming, D.W., Morris, R.V., Navarro-González, R., Niles, P.B., Owen, T., Pepin, R.O., Squyres, S., Steele, A., Stern, J.C., Summons, R.E., Sumner, D.Y., Sutter, B., Szopa, C., Teinturier, S., Trainer, M.G., Wray, J.J., Grotzinger, J.P., and the MSL Team. (2013) Volatile, isotope, and organic analysis of martian fines with the Mars Curiosity Rover. *Science* 341, doi:10.1126/science.1238937.
- Levin, G.V. (1997) The Viking Labeled Release Experiment and life on Mars. In *Proc SPIE* 3111, doi:10.1117/12.278768.
- Levin, G.V. and Straat, P.A. (1977) Recent results from the Viking Labeled Release Experiment on Mars. *J Geophys Res* 82:4663–4667.
- Lockheart, M.J., van Bergen, P.F., and Evershed, R.P. (2000) Chemotaxonomic classification of fossil leaves from the Miocene Clarkia lake deposit, Idaho, USA based on *n*-alkyl lipid distributions and principal component analyses. *Org Geochem* 31:1223–1246.
- Mahaffy, P.R., Webster, C.R., Cabane, M., Conrad, P.G., Coll, P., Atreya, S.K., Arvey, R., Barciniak, M., Benna, M., and Bleacher, L. (2012) The Sample Analysis at Mars investigation and instrument suite. *Space Sci Rev* 170:401–478.
- McAdam, A.C., Franz, H.B., Sutter, B., Archer, P.D., Freissinet, C., Eigenbrode, J.L., Ming, D.W., Atreya, S.K., Bish, D.L., Blake, D.F., Bower, H.E., Brunner, A., Buch, A., Glavin, D.P., Grotzinger, J.P., Mahaffy, P.R., McLennan, S.M., Morris, R.V., Navarro-González, R., Rampe, E.B., Squyres, S.W., Steele, A., Stern, J.C., Sumner, D.Y., and Wray, J.J. (2014) Sulfur-bearing phases detected by evolved gas analysis of the Rocknest aeolian deposit, Gale Crater, Mars. *J Geophys Res Planets* 119:373–393.
- McCollom, T.M., Ritter, G., and Simoneit, B.R. (1999) Lipid synthesis under hydrothermal conditions by Fischer-Tropsch-type reactions. *Orig Life Evol Biosph* 29:153–166.
- McKay, D.S., Gibson, E.K., Thomas-Keppta, K.L., Vali, H., Romanek, C.S., Clemett, S.J., Chilliier, X.D.F., Maechling, C.R., and Zare, R.N. (1996) Search for past life on Mars: possible relic biogenic activity in martian meteorite ALH84001. *Science* 273:924–930.
- Melendez, I., Grice, K., and Schwark, L. (2013) Exceptional preservation of Palaeozoic steroids in a diagenetic continuum. *Sci Rep* 3, doi:10.1038/srep02768.
- Milliken, R.E., Grotzinger, J.P., and Thomson, B.J. (2010) Paleoclimate of Mars as captured by the stratigraphic record in Gale Crater. *Geophys Res Lett* 37, doi:10.1029/2009gl041870.
- Ming, D.W., Mittlefehldt, D.W., Morris, R.V., Golden, D., Gellert, R., Yen, A., Clark, B.C., Squyres, S.W., Farrand, W.H., and Ruff, S.W. (2006) Geochemical and mineralogical indicators for aqueous processes in the Columbia Hills of Gusev Crater, Mars. *J Geophys Res Planets* 111, doi:10.1029/2005JE002560.
- Ming, D.W., Archer, P., Glavin, D., Eigenbrode, J., Franz, H., Sutter, B., Brunner, A., Stern, J., Freissinet, C., and McAdam, A. (2014) Volatile and organic compositions of sedimentary rocks in Yellowknife Bay, Gale Crater, Mars. *Science* 343, doi:10.1126/science.1245267.
- Moore, H.J., Hutton, R., Clow, G., and Spitzer, C., editors. (1987) *Physical Properties of the Surface Materials at the Viking Landing Sites on Mars*, US Government Printing Office, Washington, DC.
- Morris, R.V., Klingelhöfer, G., Schröder, C., Rodionov, D.S., Yen, A., Ming, D.W., de Souza, P.A., Fleischer, I., Wdowiak, T., Gellert, R., Bernhardt, B., Evlanov, E.N., Zubkov, B., Foh, J., Bonnes, U., Kankeleit, E., Gütlich, P., Renz, F., Squyres, S.W., and Arvidson, R.E. (2006) Mössbauer mineralogy of rock, soil, and dust at Gusev Crater, Mars: Spirit's journey through weakly altered olivine basalt on the plains and pervasively altered basalt in the Columbia Hills. *J Geophys Res Planets* 111, doi:10.1029/2005je002584.
- Morris, R.V., Ruff, S.W., Gellert, R., Ming, D.W., Arvidson, R.E., Clark, B.C., Golden, D., Siebach, K., Klingelhöfer, G., and Schröder, C. (2010) Identification of carbonate-rich outcrops on Mars by the Spirit rover. *Science* 329:421–424.
- Navarro-González, R. and McKay, C.P. (2011) Reply to comment by Biemann and Bada on “Reanalysis of the Viking results suggests perchlorate and organics at mid-latitudes on Mars.” *J Geophys Res Planets* 116, doi:10.1029/2011je003880.
- Navarro-González, R., Rainey, F.A., Molina, P., Bagaley, D.R., Hollen, B.J., de la Rosa, J., Small, A.M., Quinn, R.C., Grunthaner, F.J., Caceres, L., Gomez-Silva, B., and McKay, C.P. (2003) Mars-like soils in the Atacama Desert, Chile, and the dry limit of microbial life. *Science* 302:1018–1021.
- Navarro-González, R., Navarro, K.F., de la Rosa, J., Iniguez, E., Molina, P., Miranda, L.D., Morales, P., Cienfuegos, E., Coll, P., Raulin, F., Amils, R., and McKay, C.P. (2006) The limitations on organic detection in Mars-like soils by thermal volatilization–gas chromatography–MS and their implications for the Viking results. *Proc Natl Acad Sci USA* 103:16089–16094.
- Navarro-González, R., Vargas, E., de la Rosa, J., Raga, A.C., and McKay, C.P. (2010) Reanalysis of the Viking results suggests perchlorate and organics at midlatitudes on Mars. *J Geophys Res Planets* 115, doi:10.1029/2010JE003599.
- Navarro-González, R., Vargas, E., de la Rosa, J., Raga, A.C., and McKay, C.P. (2011) Correction to “Reanalysis of the Viking results suggests perchlorate and organics at midlatitudes on Mars.” *J Geophys Res Planets* 116, doi:10.1029/2011JE003854.
- Nejbert, K., Krajewski, K.P., Dubinska, E., and Pecskey, Z. (2011) Dolerites of Svalbard, north-west Barents Sea Shelf: age, tectonic setting and significance for geotectonic interpretation of the High-Arctic Large Igneous Province. *Polar Res* 30, doi:10.3402/polar.v30i0.730.
- Pavlov, A.A., Vasilyev, G., Ostryakov, V.M., Pavlov, A.K., and Mahaffy, P. (2012) Degradation of the organic molecules in the shallow subsurface of Mars due to irradiation by cosmic rays. *Geophys Res Lett* 39, doi:10.1029/2012gl052166.
- Peters, K.E., Walters, C.C., and Moldowan, J.M., editors. (2005) *The Biomarker Guide*, Cambridge University Press, Cambridge, UK.
- Petsch, S.T. (2011) Carbon (organic, degradation). In *Encyclopedia of Geobiology*, Vol. 1, edited by J. Reitner and V. Thiel, Springer, Dordrecht, the Netherlands, pp 234–238.
- Rosenbauer, H., Fuselier, S.A., Ghielmetti, A., Greenberg, J.M., Goesmann, F., Ulamec, S., Israel, G., Livi, S., MacDermott, J.A., Matsuo, T., Pillinger, C.T., Raulin, F., Roll, R., and

- Thiemann, W. (1999) The COSAC experiment on the lander of the ROSETTA mission. *Adv Space Res* 23:333–340.
- Rothschild, L.J. (1990) Earth analogs for martian life. Microbes in evaporites, a new model system for life on Mars. *Icarus* 88:246–260.
- Rushdi, A.I. and Simoneit, B.R. (2001) Lipid formation by aqueous Fischer-Tropsch-type synthesis over a temperature range of 100 to 400 degrees C. *Orig Life Evol Biosph* 31:103–118.
- Sawyer, D.J., McGehee, M.D., Canepa, J., and Moore, C.B. (2000) Water soluble ions in the Nakhla martian meteorite. *Meteorit Planet Sci* 35:743–747.
- Simmonds, P.G. (1970) Whole microorganisms studied by pyrolysis–gas chromatography–mass spectrometry: significance for extraterrestrial life detection experiments. *Appl Microbiol* 20:567–572.
- Simmonds, P.G., Shulman, G.P., and Stenbridge, C.H. (1969) Organic analysis by pyrolysis–gas chromatography–mass spectrometry a candidate experiment for the biological exploration of Mars. *J Chromatogr Sci* 7:36–41.
- Simmonds, P.G., Medley, E.E., Ratcliff, M.A., and Shulman, G.P. (1972) Thermal decomposition of aliphatic monoaminomonocarboxylic acids. *Anal Chem* 44:2060–2066.
- Squyres, S.W., Grotzinger, J.P., Arvidson, R.E., Bell, J.F., Calvin, W., Christensen, P.R., Clark, B.C., Crisp, J.A., Farand, W.H., Herkenhoff, K.E., Johnson, J.R., Klingelhöfer, G., Knoll, A.H., McLennan, S.M., McSween, H.Y., Morris, R.V., Rice, J.W., Rieder, R., and Soderblom, L.A. (2004) *In situ* evidence for an ancient aqueous environment at Meridiani Planum, Mars. *Science* 306:1709–1714.
- Steele, A., Amundsen, H.E.F., and Botta, O. (2006) The Arctic Mars Analogue Svalbard Expedition 2006. In *Mars2030—AustroMars Science Workshop*, Austrian Space Forum, Innsbruck.
- Steele, A., Fries, M.D., Amundsen, H.E.F., Mysen, B.O., Fogel, M.L., Schweizer, M., and Bocktor, N.Z. (2007) Comprehensive imaging and Raman spectroscopy of carbonate globules from martian meteorite ALH 84001 and a terrestrial analogue from Svalbard. *Meteorit Planet Sci* 42:1549–1566.
- Steele, A., McCubbin, F.M., Fries, M., Kater, L., Bocktor, N.Z., Fogel, M.L., Conrad, P.G., Glamoclija, M., Spencer, M., Morrow, A.L., Hammond, M.R., Zare, R.N., Vicenzi, E.P., Siljeström, S., Bowden, R., Herd, C.D.K., Mysen, B.O., Shirey, S.B., Amundsen, H.E.F., Treiman, A.H., Bullock, E.S., and Jull, A.J.T. (2012a) A reduced organic carbon component in martian basalts. *Science* 337:212–215.
- Steele, A., McCubbin, F.M., Fries, M.D., Golden, D.C., Ming, D.W., and Benning, L.G. (2012b) Graphite in the martian meteorite Allan Hills 84001. *Am Mineral* 97:1256–1259.
- Steininger, H., Goesmann, F., and Goetz, W. (2012) Influence of magnesium perchlorate on the pyrolysis of organic compounds in Mars analogue soils. *Planet Space Sci* 71:9–17.
- Stephan, T., Jessberger, E.K., Heiss, C.H., and Rost, D. (2003) TOF-SIMS analysis of polycyclic aromatic hydrocarbons in Allan Hills 84001. *Meteorit Planet Sci* 38:109–116.
- Taylor-George, S., Palmer, F., Staley, J.T., Borns, D.J., Curtiss, B., and Adams, J.B. (1983) Fungi and bacteria involved in desert varnish formation. *Microb Ecol* 9:227–245.
- Thiel, V. and Sjövall, P. (2011) Using time-of-flight secondary ion mass spectrometry to study biomarkers. *Annu Rev Earth Planet Sci* 39:125–156.
- Treiman, A.H. (2004) Submicron magnetite grains and carbon compounds in martian meteorite ALH84001: inorganic, abiotic formation by shock and thermal metamorphism. *Astrobiology* 3:369–392.
- Vaniman, D., Bish, D., Ming, D., Bristow, T., Morris, R., Blake, D., Chipera, S., Morrison, S., Treiman, A., and Rampe, E. (2014) Mineralogy of a mudstone at Yellowknife Bay, Gale Crater, Mars. *Science* 343, doi:10.1126/science.1243480.
- Wilson, M.A., Philp, R.P., Gillam, A.H., Gilbert, T.D., and Tate, K.R. (1983) Comparison of the structures of humic substances from aquatic and terrestrial sources by pyrolysis gas chromatography–mass spectrometry. *Geochim Cosmochim Acta* 47:497–502.
- Wright, I.P., Grady, M.M., and Pillinger, C.T. (1989) Organic materials in a martian meteorite. *Nature* 340:220–222.
- Zolotov, M.Y. and Shock, E.L. (2000) An abiotic origin for hydrocarbons in the Allan Hills 84001 martian meteorite through cooling of magmatic and impact-generated gases. *Meteorit Planet Sci* 35:629–638.

Address correspondence to:

S. Siljeström

Department of Chemistry, Materials and Surfaces

SP Technical Research Institute of Sweden

Box 587

501 15 Borås

Sweden

E-mail: sandra.siljeström@sp.se

Submitted 28 June 2014

Accepted 14 August 2014